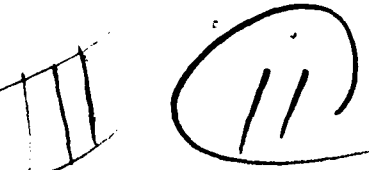


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THE EFFECTS OF AEROSOL COMPLEX  
REFRACTIVE INDEX AND SIZE DISTRIBUTION  
VARIATIONS ON EXTINCTION  
AND ABSORPTION COEFFICIENTS  
PART II: ANALYSIS OF THE  
COMPUTATIONAL RESULTS**

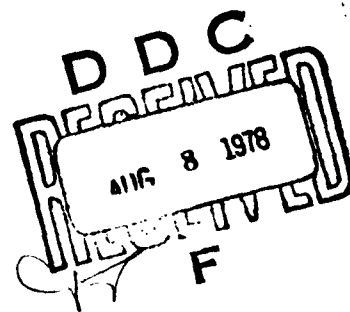
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**MARCH 1978**

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**By**

**S.G. Jennings  
J.B. Gillespie**



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**US Army Electronics Research and Development Command  
Atmospheric Sciences Laboratory  
White Sands Missile Range, N.M. 88002**

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20. ABSTRACT (Continue on reverse side if necessary and identify by block number) A comprehensive study has been made of the effect of the complex index of refraction and particle size distribution on the volume extinction and absorption coefficients. Calculations were performed by use of a generalized Mie scattering program for lognormal particle size distributions over a range of geometric mean radius $r_g$ from 0.005 $\mu m$ to 10 $\mu m$ and geometric standard deviation $\sigma_g$ from 1.0 up to 2.5. The results are normalized to a total particle mass loading of 100 $g/m^3$ and a particle density of 2.5 $g/cm^3$ . The computations were made at		

## 20. ABSTRACT (cont)

wavelength  $\lambda$  of 0.55, 0.6943, 1.06, 3.8 and 10.6 micrometers, for a range of imaginary index  $k$  from 0.001 to 6.0 and a range of real index  $n$  from 0.1 up to 6.5.

The extinction  $\sigma_E$  is insensitive to imaginary index  $k$  for  $0.001 \leq k \leq 0.025$  providing  $0.01 < r_g/\lambda \leq 0.2$ ,  $\sigma_g \geq 2.0$  and  $\lambda \leq 1.06\mu\text{m}$ , while the ratio of the extinction at real index 1.8 compared to the value at  $n = 1.33$  varies by up to a factor of about 4.0 for identical parameters. Extinction becomes nearly independent of both real index  $n$  and imaginary index  $k$  for  $r_g/\lambda \leq 0.2$ ,  $\sigma_g \geq 1.5$ , which represents most measured atmospheric size distributions for  $0.001 \leq k \leq 2.0$  and  $1.4 \leq n \leq 2.5$  providing  $0.55\mu\text{m} \leq \lambda \leq 1.06\mu\text{m}$ . The absorption  $\sigma_A$  is insensitive to  $n$  for  $0.01 \leq r_g/\lambda \leq 10$ , while  $\sigma_A$  varies linearly with  $k$  providing  $0.001 \leq k \leq 0.01$ ,  $r_g/\lambda \leq 0.5$  and  $\sigma_g \leq 2.0$ .

Absorption is independent (within a factor of approximately 3) of  $r_g$  for  $r_g/\lambda \geq 0.75$  with  $k \leq 0.05$  which represents a reasonably wide range of realistic atmospheric values. The extinction  $\sigma_E$  decreases as  $1/r_g$  for  $r_g/\lambda \geq 0.5$  and  $\geq 0.25$  for the spectral ranges  $0.55\mu\text{m} \leq \lambda \leq 1.06\mu\text{m}$  and  $3.8\mu\text{m} \leq \lambda \leq 10.6\mu\text{m}$ , respectively, for  $k \leq 0.5$ .

The real index  $n$  significantly affects  $\sigma_E$  and  $\sigma_A$  for  $3.8\mu\text{m} \leq \lambda \leq 10.6\mu\text{m}$  and extinction is mainly dependent on  $n$  over a wide range of  $r_g$  and  $\sigma_g$ . The extinction

$\sigma_E$  is well approximated by  $\sigma_E = \sigma_{\max} \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(n/n_g)}{\ln \sigma_n} \right]^2 \right\}$  where

$\sigma_{\max}$  is the extinction at  $n = n_g$ , and  $\sigma_n$  is the geometric deviation of  $\sigma_E$  with reference to  $n$ .

Extinction is invariant to  $k$  for  $0.001 \leq k \leq 8.0$  for broad realistic atmospheric distributions in the wavelength range  $3.8\mu\text{m}$  to  $10.6\mu\text{m}$ . Extinction is more sensitive to changes in real index than to imaginary index (up to a factor of about 3) for the majority of atmospheric particle size distributions in the same wavelength range.

The particle size distribution is of major importance in affecting  $\sigma_E$  and  $\sigma_A$  and has greatest influence for  $k \leq 0.1$ . Extinction is constant for  $r_g/\lambda \leq 0.15$  for  $k = 0.5$ , for  $3.8\mu\text{m} \leq \lambda \leq 10.6\mu\text{m}$ . The distribution spread  $\sigma_g$  does not appreciably influence the extinction or absorption for realistic values of  $r_g \geq 0.25\mu\text{m}$  and  $0.005\mu\text{m}$ , respectively.

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## PREFACE

This study was undertaken to answer Army problems associated with electromagnetic propagation and electro-optic sensor systems. The results of this study apply directly to the atmospheric aerosol measurements program and to research associated with the determination of the optical constants of atmospheric particulate matter. The study supplies answers to such questions as which optical quantities should be measured and how accurately they should be determined for several important wavelengths. It also gives insight to the applicability of experimental techniques used to determine the imaginary refractive index of particulates.

The Mie Scattering Computer Code, developed by Gerry W. Grams, N.C.A.R., Boulder, was made available by R. G. Pinnick. Much appreciated help was given by Don Hoihjelle and Gilbert Fernandez in the initial stages of the computer programming work. Gratitude is accorded to Ken Webb for diligent work in reducing the computer data to graphical form. Thanks are extended to J. Lentz and K. O. White for constructive suggestions.

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MIE THEORY SENSITIVITY STUDIES  
THE EFFECTS OF AEROSOL COMPLEX REFRACTIVE INDEX AND SIZE DIS-  
TRIBUTION VARIATIONS ON EXTINCTION AND ABSORPTION COEFFICIENTS  
PART II: ANALYSIS OF THE COMPUTATIONAL RESULTS

INTRODUCTION

The importance of the effects of atmospheric aerosols on the operation of electro-optical devices has increased considerably during recent years. In addition, aerosols affect radiative transfer balance in the atmosphere and hence climate. An important contribution to the heat budget of the atmosphere is attributed to absorbed radiation by polydispersions of aerosols.

The usual procedure for estimating these effects is to assume that the aerosol is composed of homogeneous spherical particles and use Mie theory [1] to calculate the scattering parameters for radiation attenuation, discernibility of targets, etc. The input parameters required for Mie calculations are (1) particle size distribution, (2) particle number density, and (3) particle refractive index as a function of radiation wavelength. Since aerosol size distribution and refractive index are in practice difficult quantities to measure, a theoretical study to assess the relative importance of these quantities in radiation calculations is imperative. In this report the authors investigate the effects of particle size distribution and real and imaginary parts of the refractive index on volume extinction and absorption coefficients for families of size distributions and ranges of refractive index that are reasonable for aerosols in the atmosphere. This theoretical sensitivity study has been performed



for selected wavelengths from the visible (0.55 $\mu\text{m}$ ) to the middle-infrared (10.6 $\mu\text{m}$ ). Effects of polydispersions of aerosol containing particles of mixed composition and irregular particles have not been addressed.

#### COMPUTATIONAL DETAILS

A generalized Mie code developed by G. W. Grams of the National Centre for Atmospheric Research was used. The Mie theoretical formalisms have already been described by several workers [2-4] and will not be further discussed here. The program was verified with a MIE-2 program of Blättner [5] and with that of Grams [6] using a different computer. The results of the calculations are presented in the form of volume extinction coefficient  $\sigma_E$  ( $\text{km}^{-1}$ ), volume scattering coefficient  $\sigma_S$  ( $\text{km}^{-1}$ ) and volume absorption coefficient  $\sigma_A$  ( $\text{km}^{-1}$ ) where, for example,  $\sigma_E$  is given by

$$\sigma_E = \int_{\tau_1}^{\tau_2} \pi r^2 Q_{\text{ext}}(x, m) n(r) dr, \quad (1)$$

where  $n(r)$  is the particle differential size distribution,  $Q_{\text{ext}}$  is the efficiency factor for extinction and is a function of the particle complex index of refraction  $m$  and the size parameter  $2\pi r/\lambda$ . Here, the complex index of refraction  $m = n - ik$ , where  $n$  is the real part of the refractive index and  $k$  the imaginary (absorptive) part. The particle radius and radiation wavelength are denoted by  $r$  and  $\lambda$ , respectively. All computations are performed with the number of particles per logarithmic radius interval,  $dn/d \ln r$  given by

$$\frac{dN}{d \ln r} = \frac{1}{\sqrt{2\pi} \ln \sigma_g} \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} \quad (2)$$

$r_g$  and  $\sigma_g$  are defined as the geometric mean radius and geometric standard deviation, where  $N$  is the total number of particles per unit volume.

A lognormal distribution was chosen to represent the particle size distribution for several reasons:

1. It represents desert-type aerosols adequately [7].
2. It also represents urban-particle size measurements [8].
3. Davies [9] has shown that particle size distributions can be well represented by combinations of lognormally distributed values of particle size.

Calculations have been made over a wide range of geometric mean radii,  $r_g$ , from  $0.005\mu\text{m}$  up to  $10.0\mu\text{m}$  and geometric standard deviation,  $\sigma_g$ , from 1.0 up to 2.5. The results have been normalized arbitrarily to a total particle mass of  $100\mu\text{g m}^{-3}$  with a uniform particle density of  $2.5\text{ g cm}^{-3}$ , consistent with measurements of maritime and dry urban aerosols by Hanel [10].

APPROXIMATION FOR THE REGIMES  $r/\lambda \ll 1$ ;  $|mx| < 1$ ,  $r/\lambda \gg 1$

Regime  $r/\lambda \ll 1$ ,  $|mx| < 1$

This is the regime for particles small compared to the wavelength and for small phase shifts. A small phase shift implies that the phase difference between the incident and scattered radiation fields is small. van de Hulst [2] has shown that the efficiency factor of absorption,  $Q_A$ , can be expressed as the series expansion

$$Q_A = -\text{Imag} \left[ 4x \left( \frac{m^2 - 1}{m^2 + 2} \right) + \frac{4}{15} x^3 \left( \frac{m^2 - 1}{m^2 + 2} \right)^2 \cdot \frac{m^4 + 27m^2 + 38}{2m^2 + 3} \right], \quad (3)$$

and  $Q_S$ , the efficiency factor for scattering can be written as

$$Q_S = x^4 \cdot 8/3 \text{Re} \left( \frac{m^2 - 1}{m^2 + 2} \right)^2, \quad (4)$$

where  $x = 2\pi r/\lambda$  and  $m = n - ik$ .

For the small particle and phase shift limit, ( $r \ll \lambda$ ),  $|mx| < 1$ , Eq. (3) becomes

$$Q_{\text{ext}} = -4 \times \text{Im} \left( \frac{m^2 - 1}{m^2 + 2} \right). \quad (5)$$

This can be written as

$$Q_A = \frac{2\pi r}{\lambda} \cdot \frac{24 nk}{(n^2 - k^2 + 2)^2 + (2nk)^2}. \quad (6)$$

The absorption coefficient,  $\sigma_a$ , for a particle of radius  $r$  can be written as

$$\sigma_a = \pi r^2 Q_A. \quad (7)$$

Therefore, the absorption per unit mass in Eq. (7) can be written as

$$\frac{\sigma_a}{m} = \sigma_A = \frac{\pi r^2}{4\pi/3 r^3 \rho} \cdot Q_A = \frac{3}{4\rho r} Q_A, \quad (8)$$

where  $\rho$  is the particle density.

Therefore, the absorption coefficient per unit mass,  $\sigma_A$ , for small particles and small phase shifts can be written as

$$\sigma_A = \frac{3}{4\rho} \frac{2\pi}{\lambda} \cdot \frac{24 nk}{(n^2 - k^2 + 2)^2 + (2nk)^2}. \quad (9)$$

Various relationships between  $\sigma_A$  and the optical constants  $n$  and  $k$  are now established below for  $r/\lambda \ll 1$ ,  $|mx| \ll 1$ . The absorption,  $\sigma_A$ , can be represented by the expression

$$\sigma_A = \text{constant} \cdot \frac{nk}{(n^2 - k^2 + 2)^2 + (2nk)^2}, \quad (10)$$

where it is assumed that  $n$  and  $k$  are invariant with wavelength  $\lambda$ .

$$1. \quad n = k: \text{ Then } \sigma_A \text{ becomes } \sigma_A = \frac{k^2}{4 + 4k^4} = \frac{n^2}{4 + 4n^4} \quad (11)$$

$$2. \quad n > k. \quad \sigma_A = \text{constant} \cdot \frac{nk}{(n^2 + 2)^2}, \quad (12)$$

or proportional to  $k$  for a fixed value of  $n$ , and proportional to

$$\frac{n}{(n^2 + 2)^2} \quad (13)$$

for a fixed value of  $k$ .

$$3. \quad k > n. \quad \sigma_A = \text{constant} \cdot \frac{nk}{(2 - k^2)^2} \quad (14)$$

Therefore,  $\sigma_A$  is proportional to  $n$  for constant  $k$  and  $\sigma_A$  is proportional to

$$\frac{k}{k^4 - 2k^2 + 4} \quad (15)$$

for a fixed value of  $n$ .

Geometrical Optics Approximation,  $r \gg \lambda$

It is known by physical argument, for example the work of van de Hulst [2] and Chylek [11], that the limit of the extinction efficiency is 2 for  $r \gg \lambda$ . Therefore the volume extinction coefficient for a distribution  $n(r)$  is twice the geometrical cross-sectional area of the polydispersion of particles, with the assumption that all of the particles in the distribution satisfy the  $r/\lambda \gg 1$  condition. This approximation is discussed further in the appendix.

#### COMPLEX REFRACTIVE INDEX FOR SEVERAL ATMOSPHERIC AEROSOL CONSTITUENTS

Atmospheric aerosols normally contain a number of constituents; the dominant one (from the point of view of radiation transfer) generally varies with location. Because of the nonlinearity of Mie theory, no single set of optical constants can adequately characterize the aerosol.

For this reason, more information is required of the constituents of natural aerosols. Mixtures of particle types are probably prevalent, which renders the determination of the refractive index of the mixture difficult and probably ambiguous. Nevertheless, knowledge of a reasonably wide range of the more commonly found atmospheric constituents is available. Table 1 gives a range of minimum and maximum values of both the real and imaginary part of the complex index of refraction for ammonium sulfate, sodium chloride, water, quartz (crystalline form), sulfuric acid solutions, ammonium sulfate solutions, carbon, and atmospheric dust over relatively narrow wavelength ranges. Toon and Pollack [12] have given a comprehensive set of values of ammonium sulfate (crystalline form), and Table 1 shows their extreme values for the three wavelength bands  $0.5\mu\text{m}$  to  $1.0\mu\text{m}$ ,  $3.0\mu\text{m}$  to  $5.0\mu\text{m}$ , and  $9.0\mu\text{m}$  to  $11.0\mu\text{m}$ . Values of the indices for ammonium sulfate solutions have been taken from the work of Downing et al. [13]. Their extrapolation of the number densities of the ions present in crystalline ammonium sulfate gives good agreement with the values of Toon et al. [12].

Toon et al. [12] have presented the refractive indices for crystalline NaCl. Table 1 shows the extreme values for the three wavelength ranges. Querry et al. [14] have recently published refractive index values of aqueous solutions of NaCl (0.25 M and 0.5 M) which closely correspond to values of solutions found in the oceans.

The index values of bulk crystalline quartz (a constituent of desert aerosols) are taken from Spitzer and Kleinman [16] and possess extreme values of  $n$  and  $k$  from 0.1 up to 7.5 and from 0.015 up to 7.5

in the wavelength region from  $9\mu\text{m}$  to  $11\mu\text{m}$ . Since crystalline quartz is birefringent, calculations for the ordinary and extraordinary rays are weighted 2 to 1 following the procedure of Peterson and Weinman [17].

Palmer and Williams [18] give the optical constants of sulfuric acid solutions (which are found in the stratosphere and troposphere) for solution concentrations of 25, 38, 50, 75, 84.5, and 95 percent. Table 1 shows the values for solution concentrations of 25, 50, and 75 percent for the three wavelength ranges.

Carbon is a common constituent in the atmosphere which constitutes about 0.5 to 1.5 percent of the aerosol in desert-type aerosols and may approach up to 10 to 15 percent of the aerosol mass in urban industrial areas. A range of values for the refractive index values for carbon [19, 20] taken from the summarized values of Twitty and Weinman [21] are shown in the table.

The optical constants of montmorillonite have been determined by Toon et al. [25] who used both Kramers-Kronig and dispersion techniques over a wavelength range from  $5\mu\text{m}$  to  $40\mu\text{m}$ . Their extreme values are shown in Table 1. There is a relatively wide variation in the values of the optical constants in the  $9\mu\text{m}$  to  $11\mu\text{m}$  range.

The complex index of refraction of atmospheric dust can be determined by a number of methods, and Fischer [26, 27], Lindberg and Laude [28], Hanel [10], Herman et al. [29], Grams et al. [30], Patterson et al. [31], and others have reported measurements. Some selected values are given in Table 2.

Few values of imaginary index of refraction exist for the wavelength ranges  $3\mu\text{m}$  to  $5\mu\text{m}$  and  $9\mu\text{m}$  to  $11\mu\text{m}$ . Fischer [27] has derived values of the mass absorption index  $k/\rho$  in these ranges, but values of the aerosol particle density must be assumed.

Few measurements are available for the real index of refraction for atmospheric dust in the middle infrared regime. Soil-derived aerosols are composed of a majority of clay minerals as shown by Hoidale and Blanco [32]. The refractive indices of these clay minerals in the  $3\mu\text{m}$  to  $5\mu\text{m}$  and  $9\mu\text{m}$  to  $11\mu\text{m}$  range are unknown.

A summary from Tables 1 and 2 of the range of complex refractive index values over these wavelength ranges of  $0.55\mu\text{m}$  to  $1.06\mu\text{m}$ ,  $3.0\mu\text{m}$  to  $5.0\mu\text{m}$ , and  $9.0\mu\text{m}$  to  $11.0\mu\text{m}$  is shown in Table 3. The extreme values for the  $9.0\mu\text{m}$  to  $11.0\mu\text{m}$  wavelength range are due to quartz, while the upper values in the two remaining wavelength bands are due to carbon.

#### PARTICLE SIZE DISTRIBUTION PARAMETERS FOR MEASURED POLYDISPERSIONS OF AEROSOLS

Lognormal parameters for measured distributions of urban, maritime, continental, desert aerosols, and fog are shown in Table 4. Most of the distributions possess values of geometric mean radius and standard geometric deviation between about  $0.1\mu\text{m}$  to  $5.0\mu\text{m}$  and from about  $1.5\mu\text{m}$  to  $2.5\mu\text{m}$ , respectively. For lognormal components with  $r_g \ll 0.1\mu\text{m}$ , the contribution of the distributions to extinction can usually be regarded as negligible because of their minimal mass loading. The role of sub-micrometer particles on the scattering parameters will be discussed later in the report.

## COMPUTATIONAL RESULTS

The computational results are presented in tabulated form by Jennings [32]. These tabulated results of volume extinction coefficient  $\sigma_E$ , are now analyzed in the form of graphical plots and tables. For the sake of brevity, only a relatively small cross-section of the results is presented. However, the results in this cross section are representative. The results are divided into two main categories:

Wavelength range 0.55 $\mu$ m to 1.06 $\mu$ m

Wavelength range 3.8 $\mu$ m to 10.6 $\mu$ m

Wavelength Range 0.55 $\mu$ m to 1.06 $\mu$ m

Volume Extinction Coefficient and Volume Absorption Coefficient as a Function of Real Index of Refraction. Values of volume extinction and volume absorption coefficients for a particle mass loading of 100 $\mu$ g m<sup>-3</sup> are shown in Figures 1 and 2 for  $\lambda = 0.55\mu$ m, and in Figure 3 for  $\lambda = 1.06\mu$ m. The real index of refraction  $n$  ranges from 1.0 to 2.5 for the imaginary index  $k = 0.01$ . Table 5 gives the range of geometric mean radius  $r_g$  and geometric standard deviation  $\sigma_g$  used in the above figures.

The following conclusions are drawn from the results:

1. The ratio of  $\sigma_E(n = 1.8)/\sigma_E(n = 1.33)$  for  $r_g/\lambda \leq 0.02$  and  $\sigma_g \leq 1.5$  is reduced to about 0.75 as shown in Table 6. The maximum value of  $\sigma_E(n = 1.8)/\sigma_E(n = 1.33)$  is about 5.0 for the larger  $r_g/\lambda$  values. Also shown are values of the ratio  $\sigma_E(n_1)/\sigma_E(n_2)$ , where  $n_1, n_2$  decrease with increasing values of imaginary index  $k$  for the same size distribution.

2. The extinction coefficient,  $\sigma_E$ , becomes almost independent of the real index  $n$  over the range  $1.3 \leq n \leq 2.5$  for values of  $r_g/\lambda \geq 0.5$



(see Figures 2 and 3). The dependence of  $\sigma_E$  on  $n$  becomes less for increasing values of  $\sigma_g$  because the large sphere approximation ( $r \gg \lambda$ ) now prevails where  $\sigma_E$  becomes independent of the optical constants  $n$  and  $k$  and solely dependent on the integrated cross-sectional area of the particles in the distribution.

3. Equations (3) and (4) show that the extinction is equal to the absorption for the condition  $r/\lambda \lesssim 0.1$ , and for  $|mx| < 1$ . However, this is no longer true as the spread ( $\sigma_g$ ) of the distribution increases as is evident from Figures 1 and 2 for  $\sigma_g = 2.5$  and Figures 1 and 3 for values of  $\sigma_g = 1.5, 2.0$ , and  $2.5$ .

4. The volume absorption coefficient is relatively insensitive to variations in the real index,  $n$ , and independent of  $\sigma_g$  over the  $r_g$  range from  $0.005$  up to  $5.0\mu\text{m}$  as illustrated in Figures 1 to 3. According to Eq. (13), the absorption  $\sigma_A$  decreases slowly with increasing  $n$ , i.e.,  $n/(n^2 + 2)^2$  for  $n \gg k$  in the small particle regime, for uniform sized particles. An increase in  $\sigma_g$  causes the absorption to increase slowly with increasing  $n$  (Figures 1 and 3). The absorption becomes independent of  $n$  for  $r_g/\lambda \gtrsim 2.0$ .

Volume Extinction Coefficient and Volume Absorption Coefficient as a Function of Imaginary Index of Refraction for  $0.55\mu\text{m} \leq \lambda \leq 1.06\mu\text{m}$ .

The extinction is nearly equal to the absorption for steeper size distributions for which  $r_g/\lambda \leq 0.1$  (Figures 4 and 6) for constant values of  $n$ . However, the extinction becomes constant with imaginary index  $k$  for increasing values of  $r_g$  and  $\sigma_g$  (Figures 4, 5, and 6). This finding is summarized in Table 7 where the range of imaginary index  $k$  over which

the extinction is nearly constant is tabulated for a wide range of  $r_g/\lambda$  and  $\sigma_g$  values. The extinction is insensitive to  $k$  for  $r_g/\lambda \gtrsim 0.2$ . This result corroborates the findings of Patterson et al. [38] that a variation in  $k$  from 0 to 0.05 had negligible effect on the comparison between measured and calculated visibilities for  $\sigma_g \approx 2.0$  and  $r_g \approx 0.5\mu\text{m}$ .

The absorption coefficient varies linearly with  $k$  for geometric mean radii  $r_g \lesssim 0.1\mu\text{m}$ , irrespective of  $\sigma_g$  for  $\sigma_g \leq 2.5$ . This is in agreement with the small particle approximation [Eq. (12)]. More generally, maximum values of  $k$  for which linearity of  $\sigma_A$  with  $k$  (within  $\pm 20$  percent) ceases for  $\lambda = 0.55\mu\text{m}$  and  $1.06\mu\text{m}$ , are shown in Tables 8 and 9 for appropriate  $r_g$  and  $\sigma_g$  values. The linearity of  $\sigma_A$  with  $k$  ceases for increasing  $r_g/\lambda$  and  $\sigma_g$  values.

The dependence of absorption on  $k$  (Figures 4, 5, and 6) is characterized by a steady increase in absorption with increasing  $k$ , then a plateau where absorption is constant, followed by a decrease in absorption for increasing values of  $k$ . The absorption becomes constant for  $k > k_c$ , and values of  $k_c$  are tabulated in Tables 10 and 11 for realistic values of  $r_g$  and  $\sigma_g$  for  $\lambda = 0.55\mu\text{m}$  and  $1.06\mu\text{m}$ . The absorption becomes insensitive for  $k$  values between 0.025 and 0.05 for most particle size distributions in this wavelength regime.

The Dependence of Extinction Coefficient and Absorption Coefficient on Particle Size for  $0.55\mu\text{m} < \lambda < 1.06\mu\text{m}$ . Bergstrom [39] discussed the extinction and absorption coefficients of the atmospheric aerosol as a function of particle size for uniform particles and for a narrow range

of power law and lognormal distributions. He used a relatively narrow range of optical constants.

Plots of  $\sigma_E$  and  $\sigma_A$  as a function of  $r_g$  and  $\sigma_g$  are presented in Figures 7 and 8, for real indices of 1.6, 1.7, and a number of imaginary index values for wavelengths 0.55 $\mu$ m and 1.06 $\mu$ m.

The absorption per unit mass is nearly independent of geometric mean radius (within a factor of 3) for  $r_g/\lambda \leq 0.75$  and  $k \leq 0.05$  in agreement with Waggoner et al. [35]. The implications of a constant response of absorption to particle size for certain conditions will be discussed in a subsequent report [36]. For  $r/\lambda \ll 1$ , Eq. (9) predicts a constant value of absorption coefficient for fixed values of  $n$ ,  $k$ , and  $\lambda$ .

The absorption per unit mass,  $\sigma_A$ , decreases as  $1/r_g$  for  $r_g/\lambda \gtrsim 7$  for  $k = 0.01$  and  $r_g/\lambda \gtrsim 0.2$  with  $k = 0.5$ . This can be seen from the following considerations: The volume absorption coefficient  $\sigma_{ABS}$  can be written for a polydispersion of particles as:

$$\sigma_{ABS} = \int_{r_1}^{r_2} \pi r^2 Q_{ABS} n(r) dr, \quad (16)$$

where  $Q_{ABS}$  is the efficiency factor for absorption. This can be rewritten as

$$\sigma_{ABS} = \sqrt{\frac{\pi}{2}} \cdot \frac{N}{\ln \sigma_g} \int_{r_1}^{r_2} Q_{ABS} r \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} dr \quad (17)$$

since

$$n(r) = \frac{N}{\sqrt{2\pi} \ln \sigma_g} \frac{1}{r} \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} \quad (18)$$

for a lognormal distribution.

Plass [42] has shown that  $Q_{ABS}$  becomes constant for  $r/\lambda > 1$  for  $k > 0.1$ , and probably for  $k > 0.05$ . Therefore, the expression for  $\sigma_A$  reduces to the form

$$\sigma_{ABS} = \text{constant} \int_{r_1}^{r_2} r \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} dr \quad (19)$$

The above integral has an identical form to the expression for the total cross-sectional area of a lognormal distribution. Information in the appendix shows that the cross-sectional area per unit volume for a lognormal distribution is proportional to  $1/r_g$  for constant  $\sigma_g$ . Therefore, the absorption per unit mass,  $\sigma_g$ , is proportional to  $1/r_g$  under the above conditions. The  $1/r_g$  relation is valid at increasing values of  $r_g/\lambda$  as  $k$  is reduced below 0.05, as is evident from Figures 7 and 8. This corroborates the finding [37] that the efficiency factor for absorption is constant for decreasing  $k$  and increasing  $r/\lambda$ .

The extinction coefficient per unit mass,  $\sigma_E$ , is sensitive to  $r_g$ ,  $\sigma_g$ , and  $k$  as shown in Figures 7 and 8. The extinction  $\sigma_E$  rises steeply with  $r_g$ , for a given  $\sigma_g$ , peaks and falls as  $1/r_g$  for  $r_g/\lambda \gtrsim 1.0$ . The  $1/r_g$  falloff for  $r_g/\lambda \gg 1$  has been shown in the appendix.

The extinction increases with increasing  $\sigma_g$ , for  $r_g/\lambda \lesssim 0.5$ , while the extinction decreases with increasing  $\sigma_g$  for  $r_g/\lambda \gtrsim 0.5$  and for  $k \leq 0.1$ . Furthermore  $\sigma_E$  is independent of  $r_g$  and  $\sigma_g$  within a factor of about 2 for  $r_g/\lambda < 0.25$  (Figures 7 and 8) for  $k = 0.5$ .

### Wavelength Range 3.8 to 10.6 Micrometers

Volume Extinction Coefficient as a Function of Real Index of Refraction  $n$  at  $\lambda = 3.8\mu\text{m}$ . The real index of refraction has a significant effect on the volume extinction coefficient  $\sigma_E$  and volume absorption coefficient  $\sigma_A$  for radiation wavelengths  $3.8\mu\text{m} \leq \lambda \leq 10.6\mu\text{m}$ . The dependence of  $\sigma_E$  and  $\sigma_A$  on  $n$  for a wide range of values of  $r_g$  and  $\sigma_g$  for constant  $k$  is shown in Figures 9 and 10.

The shape of the curves relating  $\sigma_E$  to  $n$  can in many cases be approximated to a functional form

$$\sigma_E = \sigma_{E\text{max}} \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(n/n_g)}{\ln \sigma_n} \right]^2 \right\} \quad (20)$$

where  $\sigma_{\text{max}}$  represents the real value of extinction at  $n = n_g$  and  $n$  is the geometric standard deviation of the  $n$  values centered around  $n_g$ .

Fitted points satisfying Eq. (20) are shown on  $\sigma_E$  versus  $n$  plots in Figure 10 for  $\lambda = 3.8\mu\text{m}$ . Values of the parameters  $n_g$ ,  $\sigma_n$  which give a good fit to the curves of  $\sigma_E$  versus  $n$  over most of the real index range are shown in Table 12.

The relation between extinction and real index of refraction demonstrates clearly from Figures 9 and 10 that: (a) real index  $n$  plays a significant role in effecting extinction for a particular size distribution and (b) the particle size distribution is also of major importance in determining the shape of the functional dependence of extinction on real refractive index. The extinction increases with increasing  $n$  with the exception of the regime  $r_g/\lambda < \text{about } 0.015$  when the extinction de-

creases according to Eq. (13).

Table 13 gives the range of imaginary index  $k$  for which the volume extinction coefficient  $\sigma_E$  is constant within about 15 percent, for  $\lambda = 3.8\mu\text{m}$ . It is noteworthy that for constant  $n$  (equal to 1.7 here),  $\sigma_E$  is relatively unchanged by variation in  $k$  for  $r_g \geq 0.25$  and  $\sigma_g \geq 1.5$ . The extinction also remains invariant to values of  $k$  provided  $k \lesssim 0.05$  for the radius range  $0.05 \leq r_g \leq 0.25$ , for relatively broad ( $\sigma \geq 2.0$ ) distributions. These results lead to the following conclusions:

1. The extinction per unit mass is predominantly dependent on the real index  $n$  for values of  $k$  given in Table 13. This dependence of  $\sigma_E$  on  $n$  can be approximated closely by Eq. (20).
2. In the regime of  $k$  where  $\sigma_E$  is linearly proportional to  $k$  (values given in Table 11),  $\sigma_E$  can be written as:

$$\sigma_E = c_2 k \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(n/n_g)}{\ln \sigma_n} \right]^2 \right\}, \quad (21)$$

where  $c_2 k$  is equal to  $\sigma_{E\text{max}}$ . Hence  $c_2$  is easily calculable for a specific  $k$ .

The increase in extinction at real index  $= n$ , compared to  $n = 1.4$  (real index of water at  $3.8\mu\text{m} = 1.364$ ) is shown in Table 14 as a function of  $r_g$  and  $\sigma_g$  for  $k = 0.025$ . The maximum variation of  $\sigma_E(n)/\sigma_E(n = 1.4)$  is about 2 for  $r_g \leq 0.05$  and  $\sigma_g \leq 2.0$ , but possesses values up to 7 for  $r_g \geq 0.1$  with  $\sigma_g \geq 1.5$ .

The range of  $k$  for which  $\sigma_E$  is nearly linear with  $k$  at a selection of  $r_g$  and  $\sigma_g$  values is given in Table 15. The linearity of  $\sigma_E$  with  $k$  only

prevails for the smaller particle range ( $\lesssim 0.1\mu\text{m}$ ).

Comment on the volume absorption coefficient as a function of  $k$ , although plotted in Figure 11 for the sake of completeness, will be deferred to a subsequent report [41].

Volume Extinction Coefficient as a Function of Real Index of Refraction at  $\lambda = 10.6\mu\text{m}$ . Figures 12 to 14 show the effect of real index of refraction  $n$  on extinction,  $\sigma_E$ , for a wide range of particle size distribution parameters  $r_g$  and  $\sigma_g$ . The range of particle distribution parameters used in the above figures is shown in Table 16. The variation of extinction with  $n$  can be represented, for most combinations of  $r_g$  and  $\sigma_g$ , by an equation similar to Eq. (21), i.e.,

$$\sigma_E = a_1 \exp [-b_1 \ln^2 (n/n_g)] \quad (22)$$

where  $a_1$  is the maximum extinction value for a given extinction versus real index curve and  $b_1$  is equal to  $1/2 \cdot \ln \sigma_n^2$  where  $\sigma_n$  is a measure of the spread of the  $\sigma_E$  values about the geometric mean  $n_g$ .

A set of points satisfying Eq. (22) is plotted along with rigorous Mie results in Figures 15 and 16 for various values of  $r_g$  and  $\sigma_g$ .

The parameters of Eq. (22) which give a reasonably good fit to the Mie results over most of the real refractive index range are given in Table 17.

Table 18 presents a summary of computations made of extinction for real refractive index  $n$ , ranging between 1.2 and 3.5 compared to values

at  $n = 1.2$  (real index of water at  $10.6\mu\text{m} = 1.185$ ), over a wide range of  $r_g$  values. The ratio is sensitive to  $\sigma_g$  for any given  $r_g$  except for  $r_g \leq 0.01$ . Table 18 represents most combinations of typical parameters of particle size distributions for atmospheric aerosols. The ratio of  $\sigma_E$  ( $n = 2.5$ )/ $\sigma_E$  ( $n = 1.2$ ) attains a maximum value of 11.6 and decreases to a value of about 2.0 at  $r_g = 2.5\mu\text{m}$ ,  $\sigma_g = 2.0$ . In general, for  $r_g \leq 0.75\mu\text{m}$  the extinction ratio increases with  $\sigma_g$  up to  $\sigma_g \leq 2.0$  and thereafter decreases with increasing  $\sigma_g$ .

Volume Extinction Coefficient as a Function of Imaginary Index of Refraction at  $\lambda = 10.6\mu\text{m}$ . The variation of  $\sigma_E$  with  $k$ , for  $n = 1.8$  is shown in Figure 17 for  $\lambda = 10.6\mu\text{m}$ . The range of imaginary index  $k$  for which  $\sigma_E$  is nearly constant (to within  $\pm 15$  percent) is summarized in Table 19. The extinction is nearly independent (within  $\pm 15$  percent) of  $k$ ,  $0.001 \leq k \leq 0.1$  for broad distributions with  $r_g = 0.25\mu\text{m}$ ,  $0.5\mu\text{m}$ . Extinction is insensitive (to within  $\pm 15$  percent) to a wide variation in  $k$  ( $0.001 \leq k \leq 8.0$ ) for  $r_g \geq 0.75\mu\text{m}$  and  $\sigma_g \geq 2.0$ . Since most atmospheric constituents have imaginary index  $k$  values within the above range, it is concluded that at  $10.6\mu\text{m}$  wavelength, extinction is relatively independent of  $k$  for the broader particle size distributions with geometric mean radii  $\geq 0.5\mu\text{m}$ . Over the range of  $k$  for which  $\sigma_E$  is relatively constant,  $\sigma_E$  can be written solely in terms of real index  $n$  according to Eq. (20).

Table 20 shows the range of  $k$  for which  $\sigma_E$  is linear to within 15 percent, with  $k$ , at  $\lambda = 10.6\mu\text{m}$ . The extinction is written in terms of Eq. (21), analogous to the  $\lambda = 3.8\mu\text{m}$  case, for the relevant values of  $r_g$



and  $\sigma_g$  shown in Table 16. Thus, the linearity of  $\sigma_E$  with  $k$  exists for particle sizes less than about  $0.25\mu\text{m}$  for values of  $k$  typical of atmospheric particulates (from about 0.01 to about 0.5 at  $\lambda = 10.6\mu\text{m}$ ). The effect of imaginary index on absorption at  $\lambda = 10.6\mu\text{m}$  will be discussed in a subsequent report [41] which will include a comment on the use of the pellet spectrophone technique [43] as a possible method to determine the imaginary index of atmospheric particulates.

Table 21 summarizes the relative sensitivity of volume extinction coefficient  $\sigma_E$  to changes in imaginary and real index of refraction. The variation in extinction as real index is increased two-fold is compared to the variation in extinction with imaginary index. It is evident that  $\sigma_E$  is more sensitive to changes in  $n$  compared to changes in  $k$  for the majority of  $r_g$  and  $\sigma_g$  values, particularly for values of  $r_g \geq 0.25\mu\text{m}$ , where extinction becomes insensitive to  $k$ . Most components of measured atmospheric particle size distributions have  $r_g$  and  $\sigma_g$  values ranging from about 0.1 to approximately  $10\mu\text{m}$  and from a minimum of about 1.25 to about 2.5, respectively, as seen from Table 4. Since typical values of imaginary index of atmospheric constituents range from about 0.1 - 0.5 at  $\lambda = 10.6\mu\text{m}$  (see Tables 1 and 2), it is concluded that the real index causes a much larger increase (up to a factor of 3) in extinction than the imaginary index.

The Effect of Particle Size Distribution on Volume Extinction Coefficient for  $\lambda = 3.8\mu\text{m}$  and  $10.6\mu\text{m}$ . The extinction  $\sigma_E$  is plotted as a function of geometric mean radius  $r_g$  for  $\sigma_g = 1.0, 1.5$ , and  $2.0$  for  $\lambda = 3.8\mu\text{m}$  and  $10.6\mu\text{m}$  in Figures 18 and 19, respectively. The real

index was fixed at 1.8 for both wavelengths, while several realistic values of the imaginary index were chosen. The results show that extinction remains insensitive to changes in  $r_g$  up to about  $0.1\mu\text{m}$  for  $\sigma_g \leq 1.5$ . The optical important range shifts to a smaller size regime with increasing  $\sigma_g$ . The extinction decreases as  $1/r_g$  for  $r_g/\lambda \gtrsim 0.25$  with the exception of  $k = 0.5$ . The extinction per unit mass decreases as  $1/r_g$  for  $r_g/\lambda > 0.3$  at  $k = 0.5$  for  $\lambda \geq 3.8\mu\text{m}$ . This is in accordance with the predictions of the appendix calculations.

It is also noteworthy that  $r_g$  has the largest effect on  $\sigma_E$  for the lower values of  $k$ . Extinction remains constant (within a factor of about 2) with increasing values of  $r_g$  up to  $1.0\mu\text{m}$  for  $\lambda = 3.8\mu\text{m}$  and  $10.6\mu\text{m}$  for  $k = 0.5$ . Since typical particle size distributions of atmospheric aerosols which contribute significantly to extinction are in the range of  $r_g$  from about  $0.1\mu\text{m}$  up to about  $1.0\mu\text{m}$  (Table 4), the results imply that extinction is relatively independent of particle size distribution for atmospheric aerosols with large values of imaginary index ( $\sim 0.5$  or greater).

The dependence of volume absorption coefficient  $\sigma_A$  on  $r_g$  is discussed fully in a subsequent report [36].

#### DEPENDENCE OF EXTINCTION COEFFICIENT AND ABSORPTION COEFFICIENT ON WAVELENGTH AND STANDARD GEOMETRIC DEVIATION

##### The Effect of Wavelength $\lambda$ on Extinction and Absorption Coefficients

The change in  $\sigma_E$  and  $\sigma_A$  as  $\lambda$  is increased from  $0.55\mu\text{m}$  to  $10.6\mu\text{m}$  is shown in Figure 20 for selected particle size distributions, for  $k = 0.01$  and  $n = 1.7$ . The absorption coefficient  $\sigma_A$  varied as  $1/\lambda$  as predicted

by Eq. (9) for  $r_g/\lambda \leq 0.05$  and  $\sigma_g \leq 1.5$ .

The extinction  $\sigma_E$  becomes nearly constant with  $\lambda$  for  $\lambda \leq 3.8\mu\text{m}$  for  $r_g \geq 1.0\mu\text{m}$  and  $k = 0.01$ , irrespective of the  $\sigma_g$  value. Some selected values of extinction and absorption coefficients at  $\lambda = 0.55\mu\text{m}$ ,  $1.06\mu\text{m}$ , and  $10.6\mu\text{m}$  are given in Table 22 for a wide range of  $r_g$  and  $\sigma_g$  values.

The table shows that  $\sigma_E$  decreases by 2 and 3 orders of magnitude as  $\lambda$  increases from  $0.55\mu\text{m}$  to  $10.6\mu\text{m}$  for  $r_g$  up to about  $0.5\mu\text{m}$ . For  $r_g \geq 0.5\mu\text{m}$ , increasing values of  $\sigma_g$  diminish the relative decrease of both  $\sigma_E$  and  $\sigma_A$  with  $\lambda$ .

These Mie results can be approximated by a power law between  $\sigma_E$ ,  $\sigma_A$ , and  $\lambda$ :

$$\sigma_A = c\lambda^y \quad (23)$$

where the exponent  $y$  initially increases from  $-1$  to a maximum of about  $-1.7$  for  $\sigma_A$  and thereafter decreases with increasing  $\sigma_g$  and  $r_g$ . The parameters for various values of  $k$  and  $n$  are found in Table 23. The extinction exponent attains a maximum value of  $-3.5$  for particles of radius  $0.1\mu\text{m}$ , becomes positive at  $r_g \approx 0.5\mu\text{m}$ , and thereafter becomes constant with increasing  $r_g$ .

An increase in  $k$  from  $0.01$  to  $0.1$  causes a decrease in the exponent for extinction for  $r_g = 0.01\mu\text{m}$ ,  $0.1\mu\text{m}$ , and  $1.0\mu\text{m}$ , and a decrease in the exponent for absorption for  $r_g = 0.1\mu\text{m}$  and  $1.0\mu\text{m}$ . The relative insensitivity of  $\sigma_E$  and  $\sigma_A$  to  $\lambda$  at large  $r_g/\lambda$  is a consequence of the increasing dependence of the coefficients on the geometrical cross-sectional

area of the particles.

Table 24 shows a comparison of  $\sigma_E$  and  $\sigma_A$  for particular size distributions which are generally representative of soil derived aerosols [5], at  $\lambda = 1.06\mu\text{m}$ ,  $3.8\mu\text{m}$ , and  $10.6\mu\text{m}$ . Different values of  $k$  are assigned to the distributions at the three wavelengths and are considered to represent a range of realistic values of  $k$  for atmospheric aerosol at these wavelengths.

The ratio of  $\sigma_E (\lambda = 3.8)/\sigma_E (\lambda = 10.6)$  increases with decreasing  $r_g$  values and possesses values of about 10, 5, and 3 for  $r_g = 0.25$ ,  $1.0$ , and  $0.5$ . Extinction at  $\lambda = 1.06\mu\text{m}$  and  $3.8\mu\text{m}$  is nearly identical for  $r_g = 0.5\mu\text{m}$  and  $1.0\mu\text{m}$  but differs by a factor of about 2.0 at  $r_g = 0.25$ . Absorption is relatively constant for the three wavelengths at  $r_g = 0.5\mu\text{m}$  and  $1.0\mu\text{m}$  but differs by up to a factor of 5 for the lowest  $r_g$  value.

#### The Influence of Geometric Standard Deviation Volume Extinction and Absorption Coefficient

Figure 21 shows the effect of geometric standard deviation  $\sigma_g$  on  $\sigma_E$  for  $\lambda = 0.55\mu\text{m}$ ,  $1.06\mu\text{m}$ ,  $3.8\mu\text{m}$ , and  $10.6\mu\text{m}$ , and for a range of  $r_g$  from  $0.005\mu\text{m}$  up to  $0.1\mu\text{m}$ . The following conclusions can be drawn:

1.  $\sigma_g$  has the largest effect on  $\sigma_E$  for  $\lambda \leq 1.06\mu\text{m}$  and for  $r_g \leq 0.1\mu\text{m}$ .
2. Increases in  $\sigma_E$  of up to an order of magnitude or greater are noted for  $r_g \leq 0.1$ , as  $\sigma_g$  is increased from 1.0 to 2.5. For the larger particle regime,  $0.25 \leq r_g \leq 10\mu\text{m}$ , variation of  $\sigma_g$  from 1.0 to 2.0 causes a change in  $\sigma_E$  of not more than a factor of about 3 for the wavelengths  $0.55\mu\text{m}$ ,  $1.06\mu\text{m}$ ,  $3.8\mu\text{m}$ , and  $10.6\mu\text{m}$ .

## CONCLUSIONS

Wavelength Range 0.55 $\mu$ m to 1.06 $\mu$ m

The ratio of volume extinction coefficient  $\frac{\sigma_E (n = 1.8)}{\sigma_E (n = 1.33)}$  which

covers the real index range for the majority of atmospheric aerosols has values up to a factor of 5 for  $0.01 \leq r_g/\lambda \leq 0.2$  in the wavelength range 0.55 $\mu$ m to 1.06 $\mu$ m for imaginary index  $k = 0.005$ . This result is in contrast to the nearly constant value of extinction as imaginary index  $k$  is varied from 0.001 to about 0.025 at  $n = 1.5, 1.7$  over the  $r_g/\lambda$  range  $0.01 < r_g/\lambda \leq 0.2$ .

The extinction coefficient becomes relatively independent of both real and imaginary index for  $r_g/\lambda > \text{about } 0.2$  over a wide range of  $k$  (0.001 to 2.0) and  $n$  (1.4 to 2.5) over a wide range of particle size distributions. Since typical particle size distributions of particles in the atmosphere possess geometric mean radii  $r_g \gtrsim 0.01\mu\text{m}$ , the results suggest that wide variations in the optical constants over the wavelength range 0.55 $\mu$ m to 1.06 $\mu$ m do not influence the extinction.

These results emphasize the relative insensitivity of extinction to  $n$  and  $k$  for typical particle size distributions measured, and the results should be carefully considered in connection with current and possible future measurements of optical constants in the wavelength region up to

at least  $1.06\mu\text{m}$ .

The volume absorption coefficient is almost insensitive to variations in the real refractive index and independent of particle size distribution for  $0.01 \leq r_g/\lambda \leq 10$  which covers the majority of atmospheric particle size distribution values of geometric mean radius.

The absorption  $\sigma_A$  increases linearly with imaginary index for  $k$  values from 0.001 up to about 0.01 for  $r_g/\lambda \leq 0.5$  and  $\sigma_g \leq 2.0$ . This  $k$  range for linearity of  $\sigma_A$  with  $k$  is narrowed for increasing  $r_g/\lambda$  values.

Absorption per unit mass is nearly independent of particle size distribution (within a factor of about 3) for  $r_g/\lambda \leq 0.75$  and  $k \leq 0.05$ , and decreases as  $1/r_g$  for  $r_g/\lambda > 7.0$  for  $k \geq 0.05$ . Typical atmospheric values of  $k$  vary from about 0.005 to 0.10 for the wavelength range  $0.55\mu\text{m}$  to  $1.06\mu\text{m}$ .

The extinction per unit mass increases with increasing  $\sigma_g$  for  $r_g/\lambda \leq 0.5$ , while the extinction decreases with increasing  $\sigma_g$  for  $r_g/\lambda \geq 0.5$ , for  $k \leq 0.1$ . The extinction  $\sigma_E$  is independent of  $r_g$  and  $\sigma_g$  within a factor of about 2 for  $r_g/\lambda < 0.25$  for  $k = 0.5$ . In addition,  $\sigma_E \approx 1/r_g$  for  $r_g/\lambda \geq 0.5$  over the range  $0.005 \leq k \leq 0.5$ .

#### Wavelength Range $3.8\mu\text{m}$ to $10.6\mu\text{m}$

Variations in the real index of refraction  $n$  affects significantly the volume extinction coefficient  $\sigma_E$  and volume absorption coefficient  $\sigma_A$  for the wavelength range  $3.8\mu\text{m}$  to  $10.6\mu\text{m}$ . Extinction is more sensitive to  $n$  than  $k$  (up to a factor of about 3) for the majority of size

distribution parameters which well represent typically measured atmospheric aerosol values.

The extinction  $\sigma_E$  is well approximated by the following formula:

$$\sigma_E = \sigma_{E \max} \exp \left\{ -\frac{1}{2} \left[ \frac{(\ln n / n_g)}{\ln \sigma_n} \right]^2 \right\}$$

where  $\sigma_{E \max}$  is the extinction at the maximum value of  $n = n_g$ , where  $\sigma_n$  is the geometric standard deviation of  $\sigma_E$  distributed with respect to  $n$ . This formula holds over a wide range of  $r_g$  and  $\sigma_g$  for  $\lambda = 3.8\mu\text{m}$  and  $10.6\mu\text{m}$ .

In the regime where extinction is linearly proportional to  $k$ , extinction  $\sigma_E$  can be written as

$$\sigma_E = ak \exp \left\{ -\frac{1}{2} \left[ \frac{\ln (n/n_g)}{\ln \sigma_n} \right]^2 \right\},$$

where  $a$  is a constant.

Extinction is invariant to a relatively wide range of imaginary index 0.001 to 8.0 for  $r_g/\lambda \gtrsim 0.08$  and  $\sigma_g \geq 2.0$  and invariant for  $k$  over the same range for  $r_g \gtrsim 0.25\mu\text{m}$  and  $\sigma_g > 2.0$  for both  $3.8\mu\text{m}$  and  $10.6\mu\text{m}$ . The  $r_g/\lambda$  range above covers some of the measured aerosol size distribution geometric mean radii for  $3.8\mu\text{m} \leq \lambda \leq 10.6\mu\text{m}$ . Since the  $k$  range 0.001 to 0.5 covers most of the imaginary index values of atmospheric constituents, it is concluded that extinction is insensitive to imaginary index  $k$  for  $r_g$  values  $> 0.3\mu\text{m}$  and  $0.8\mu\text{m}$  for  $3.8\mu\text{m}$  and  $10.6\mu\text{m}$ .

The particle size distribution is of major importance in determining the dependence of extinction and absorption on real index  $n$ .

The geometric mean radius  $r_g$  has the largest effect on extinction for  $k \leq 0.1$ . Extinction is independent of particle size distribution, within a factor of about 2 for  $r_g/\lambda \leq 0.15$ ,  $\sigma_g \leq 2.0$ , for  $k = 0.5$ . The extinction decreases as  $1/r_g$  for  $r_g/\lambda > \sim 0.25$  for  $k \leq 0.5$ .

The extinction is reduced by factors of 10, 5, and about 3 at  $\lambda = 10.6\mu\text{m}$  compared to 3.8 for reasonably representative values of both atmospheric particle size distributions ( $r_g = 0.25, 1.0$ , and  $0.5$ , respectively) and optical constants. For a typical range of  $r_g$  in the atmosphere  $0.25 \leq r_g \leq 10\mu\text{m}$ , an increase in geometric standard deviation from 1.0 to 2.0 has relatively little influence (less than factor of about 3) on the extinction, while changes in absorption of not more than a factor of 2 are effected for  $0.005 \leq r_g \leq 10\mu\text{m}$  for the same variation in  $\sigma_g$ .



# APPENDIX

## VOLUME EXTINCTION COEFFICIENT FOR A LOGNORMAL DISTRIBUTION WITH PARTICLE RADIUS $r \gg \text{WAVELENGTH } \lambda$ .

$$n(r) = \frac{N}{\sqrt{2\pi} \ln \sigma_g} \frac{1}{r} \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} \quad (\text{A-1})$$

for a Lognormal distribution

Extinction coefficient for  $r \gg \lambda$

$$= 2 \int_0^\infty \pi r^2 n(r) dr \quad (\text{A-2})$$

$$= \frac{2\pi N}{\sqrt{2\pi} \ln \sigma_g} \int_0^\infty r \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} dr \quad (\text{A-3})$$

$$\text{Volume} = \frac{4\pi}{3} \int_0^\infty r^3 n(r) dr \quad (\text{A-4})$$

$$= \frac{4\pi}{3} \frac{N}{\sqrt{2\pi} \ln \sigma_g} \int_0^\infty r^2 \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} dr \quad (\text{A-5})$$

$$\frac{\text{Extinction Coefficient for } r \gg \lambda}{\text{Volume}} = \frac{\frac{3}{2} \int_0^\infty r \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} dr}{\int_0^\infty r^2 \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} dr} \quad (\text{A-6})$$

Examine

$$\int_0^\infty r \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(r/r_g)}{\ln \sigma_g} \right]^2 \right\} dr \quad (\text{A-7})$$

This has the form

$$\int_0^\infty r \exp(-k \ln^2 cr) dr$$

$$\text{where } k = \frac{1}{2 \ln^2 \sigma_g} \text{ and } c = \frac{1}{r_g} \quad (\text{A-8})$$

Let  $y = \ln cr$

$$\therefore r = \frac{\exp y}{c} \quad \text{and} \quad dr = \frac{\exp(y)}{c} dy \quad (\text{A-9})$$

Therefore Eq. (A-7) becomes

$$\frac{1}{c^2} \int_{-\infty}^{+\infty} \exp(2y) \exp(-ky^2) dy \quad (\text{A-10})$$

According to Gradshteyn and Ryzhik [44], integrals of the form

$$\int_{-\infty}^{+\infty} \exp(-p^2 y^2 + qy) dy$$

have a solution equal to

$$\exp\left(\frac{q^2}{4p^2}\right) \cdot \frac{\sqrt{\pi}}{p} \quad \text{for } p > 0$$

Therefore, Eq. (A-10) reduces to

$$r_g^2 \cdot \sqrt{2\pi} \ln \sigma_g \exp(2 \ln^2 \sigma_g) \quad (\text{A-11})$$

We now consider

$$\int_0^{\infty} r^2 \exp\left\{-\frac{1}{2} \left[\frac{\ln r/r_g}{\ln \sigma_g}\right]^2\right\} dr \quad (\text{A-12})$$

This expression is written as

$$\int_0^{\infty} r^2 \exp[-k \ln^2 cr] dr$$

using the same constants as in (A-8).

Eq. (A-12) reduces to

$$\frac{1}{c^3} \int_{-\infty}^{\infty} \exp(3y) \exp(-ky^2) dy \quad (\text{A-13})$$

which is reducible to the following form:

$$r_g^3 \sqrt{2\pi} \ln \sigma_g \exp(4.5 \ln^2 \sigma_g) \quad (\text{A-14})$$

using the integral solution of [39].

Therefore the expression for extinction coefficient for  $r \gg \lambda$   
volume of distribution

is equal to

$$\frac{3}{2} \cdot \frac{1}{r_g} \exp (-2.5 \ln^2 \sigma_g) \quad (A-15)$$

TABLE 1. REAL AND IMAGINARY PART OF THE COMPLEX REFRACTIVE INDEX OF SEVERAL ATMOSPHERIC CONSTITUENTS

Constituent	Refractive Index	Wavelength		
		0.55 $\mu$ m to 1.06 $\mu$ m	3.0 $\mu$ m to 5.0 $\mu$ m	9.0 $\mu$ m to 11.0 $\mu$ m
(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub> [ref 12] (crystalline form)	n* k**	1.52 to 1.55 Upper limit of 1 x 10 <sup>-7</sup>	1.27 to 1.63 0.006 to 0.33	0.99 to 2.86 0.02 to 2.1
(NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub> [ref 13] 3.2 m solution	n k	- -	1.2 to 1.5 0.0 to 0.24	1.3 to 1.65 0.06 to 0.56
NaCl [ref 12] (crystalline)	n k	1.53 to 1.55 Upper limit of 1 x 10 <sup>-7</sup>	1.52 to 1.51 Upper limit of 1 x 10 <sup>-7</sup>	1.49 to 1.50 1.0 x 10 <sup>-7</sup> to 1.9 x 10 <sup>-7</sup>
NaCl [ref 14] 0.25 m solution	n k	Not given Not given	1.325 to 1.475 0.004 to 0.268	1.156 to 1.263 0.04 to 0.093
NaCl 0.5 m solution	n k	Not given Not given	1.327 to 1.481 0.000 to 0.262	1.162 to 1.265 0.039 to 0.091
H <sub>2</sub> O [ref 15]	n k	1.327 to 1.335 1.0 x 10 <sup>-9</sup> to 3.48 x 10 <sup>-6</sup>	1.325 to 1.483 0.003 to 0.272	1.153 to 1.262 0.04 to 0.097
Quartz [ref 16] (crystalline) Ordinary ray	n k	1.539 -	1.412 to 1.48 0.00001 to 0.00079	0.11 to 6.38 0.017 to 7.51
Extraordinary ray	n k	1.584 -	Not given Not given	0.11 to 7.51 0.015 to 6.37
25% solution of H <sub>2</sub> SO <sub>4</sub> in water [ref 18]	n k	1.359 to 1.366 3.02 x 10 <sup>-8</sup> to 2.75 x 10 <sup>-6</sup>	1.324 to 1.431 0.028 to 0.193	1.346 to 1.442 0.151 to 0.273
50% solution of H <sub>2</sub> SO <sub>4</sub> in water [ref 18]	n k	1.389 to 1.397 2.07 x 10 <sup>-8</sup> to 2.09 x 10 <sup>-6</sup>	1.329 to 1.379 0.076 to 0.141	1.433 to 1.650 0.199 to 0.463
75% solution of H <sub>2</sub> SO <sub>4</sub> in water [ref 18]	n k	1.422 to 1.431 2.07 x 10 <sup>-8</sup> to 1.53 x 10 <sup>-6</sup>	1.294 to 1.405 0.099 to 9.161	1.589 to 1.947 0.271 to 0.711
Carbon [19, 20, 21, 22] coals and soots	n k	1.5 to 2.1 0.015 to 0.65	1.55 to 2.0 0.25 to 0.8	2.0 to 2.45 0.35 to 1.2
20% solution of ZnCl in water [ref 23]	n k	Not given Not given	1.365 to 1.491 0.012 to 0.256	1.207 to 1.302 0.048 to 0.095
30% solution of ZnCl in water [ref 23]	n k	Not given Not given	1.377 to 1.497 0.004 to 0.226	1.232 to 1.320 0.041 to 0.092
40% solution of ZnCl in water [ref 23]	n k	Not given Not given	1.402 to 1.506 0.010 to 0.200	1.265 to 1.344 0.042 to 0.088
Limestone [ref 24]	n k	1.550 to 1.569 0.025 to 0.052	1.467 to 1.543 0.045 to 0.068	1.545 to 1.723 0.169 to 0.201
Montmorillite [ref 25]	n k	Not evaluated Not evaluated	1.37 to 1.42 (at 5.0 $\mu$ m) 0.004 to 0.005 (at 5.0 $\mu$ m)	0.71 to 2.78  0.15 to 2.06

\*n = real index.

\*\*k = imaginary index.

TABLE 2. REAL AND IMAGINARY PART OF THE COMPLEX INDEX OF REFRACTION OF ATMOSPHERIC PARTICLES, DUST, AND MARINE AEROSOL.

Constituent	Refractive Index	Wavelength		
		0.55 $\mu$ m to 1.06 $\mu$ m	3.0 $\mu$ m to 5.0 $\mu$ m	9.0 $\mu$ m to 11.0 $\mu$ m
Urban air density = 2.7 g cm <sup>-3</sup> [ref 26]	n* k**	Not given 0.0049-0.008	Not measured Not measured	Not measured Not measured
Maritime air density = 2.4 g cm <sup>-3</sup> [ref 26]	n k	Not measured 0.00033-0.0011	Not measured Not measured	Not measured Not measured
Desert aerosol [ref 28]	n k	Not measured 0.004-0.10	Not measured Not measured	Not measured Not measured
Urban aerosol [ref 10]	n k	1.485-1.66 Not measured	Not measured Not measured	Not measured Not measured
Urban aerosol density = 2.7 g cm <sup>-3</sup> [ref 27]	n k	Not measured Not measured	No measurement 0.104-0.131	Not measured 0.139-0.365
Marine aerosol density = 2.4 gm <sup>-3</sup> [ref 27]	n k	Not measured Not measured	No measurement 0.075-0.086	Not measured 0.052-0.081
Desert aerosol assumed density 2.5 gm <sup>-3</sup> [ref 27]	n k	Not measured Not measured	No measurement 0.028-0.13	Not measured 0.0025-0.375

\* n = real index

\*\* k = imaginary index

TABLE 3. RANGE OF MEASURED OPTICAL CONSTANTS OF  
SEVERAL ATMOSPHERIC CONSTITUENTS

(Ammonium Sulfate, Sodium Chloride, Water,  
Sulfuric Acid Solutions, Acid and Oil  
Smokes, Carbon, Quartz, and Atmospheric  
Dust)

<u>Wavelength Range</u> <u>(<math>\mu\text{m}</math>)</u>	<u>Real Index</u>	<u>Imaginary Index</u>
0.55 to 1.0	1.32 to 2.1	0 to 0.8
3.0 to 5.0	1.2 to 3.1	0 to 1.35
9.0 to 11.0	0.11 to 7.51	0 to 7.51

TABLE 4. TYPICAL LOGNORMAL PARAMETERS FOR  
MEASURED ATMOSPHERIC AEROSOL  
POLYDISPERSIONS

<u>Types of Atmospheric Aerosol</u>	<u>Lognormal Component</u>	<u>Geometric Mean Radius (<math>\mu\text{m}</math>)</u>	<u>Geometric Standard Deviation</u>
Heavy aerosol (mass $\approx 10^4 \mu\text{gm}^{-3}$ ) [ref 7]	1	0.47	2.2
	2	26.0	1.37
Moderate aerosol (mass $\approx 10^3 \mu\text{gm}^{-3}$ ) [ref 7]	1	0.12	2.41
	2	8.0	2.0
Light aerosol (mass $\approx 10^2 \mu\text{gm}^{-3}$ ) [ref 33]	1	0.03	2.4
	2	1.6	1.6
Fog (mass $\approx 5 \times 10^5 \mu\text{gm}^{-3}$ ) [ref 34]	1	0.25	2.1
	2	4.5	1.65
Heavy Fog (mass $\approx 10^6 \mu\text{gm}^{-3}$ ) [ref 34]	1	0.5	2.5
	2	4.5	1.8
Aerosol measured from aircraft altitude $\leq 5$ km (mass $\approx 50 \mu\text{gm}^{-3}$ ) [ref 35]	1	0.375	1.25
	2	0.77	1.5
	3	2.85	1.25
	4	3.75	1.5
Maritime aerosol (mass $\approx 30 \mu\text{gm}^{-3}$ ) [ref 36]	1	0.05	1.4
	2	0.133	1.4
	3	0.44	1.7
	4	1.50	1.7
	5	9.0	1.6

TABLE 5. PARTICLE SIZE DISTRIBUTION PARAMETERS USED  
IN THE PLOTS OF EXTINCTION AND ABSORPTION  
COEFFICIENTS, WITH REAL INDEX  $n$  (FIGURES  
1 TO 8) FOR IMAGINARY INDEX  $k = 0.01$

Wavelength $\lambda$ ( $\mu\text{m}$ )	Geometric Mean Radius $r_g$ ( $\mu\text{m}$ )	Geometric Standard Deviation $\sigma_g$
0.55	0.005	1.0
0.55	0.005	2.0
0.55	0.005	2.5
0.55, 1.06	0.01	1.25
0.55, 1.06	0.01	2.0
0.55, 1.06	0.01	2.5
0.55, 1.06	0.05	1.0
0.55, 1.06	0.05	1.5
0.55, 1.06	0.05	2.0
0.55, 1.06	0.05	2.5
0.55, 1.06	0.1	1.0
0.55, 1.06	0.1	1.5
0.55, 1.06	0.1	2.0
0.55	0.25	1.0
0.55	0.25	1.5
0.55	0.25	2.0
0.55, 1.06	0.5	1.0
0.55, 1.06	0.5	2.0
0.55	0.75	1.0
0.55	0.75	1.5
0.55, 1.06	1.0	1.0
0.05, 1.06	1.0	1.5
0.55	2.5	1.0
0.55	2.5	1.5
0.55, 1.06	5.0	1.0
0.55, 1.06	5.0	1.5



TABLE 6. THE RATIO OF EXTINCTION COEFFICIENTS  $\sigma_E(n)/\sigma_E(n=1.33)$   
AS A FUNCTION OF  $r_g/\lambda$  AND  $\sigma_g$ ; IMAGINARY INDEX  $k = 0.005$

$n/\sigma_g$	$\sigma_E(n)/\sigma_E(n = 1.33)$													
	$r_g/\lambda = 0.01$			$r_g/\lambda = 0.02$			$r_g/\lambda = 0.1$			$r_g/\lambda = 0.2$				
	1.0	1.5*	2.0*	2.5*	1.125	2.0*	2.5*	1.0	1.5*	2.0*	2.5*	1.0	1.5*	2.0*
1.33	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0
1.4	0.955	0.959	1.058	1.389	0.962	1.284	1.42	1.259	1.410	1.349	1.130	1.426	1.458	1.169
1.5	0.890	0.902	1.161	2.086	0.913	1.77	2.11	1.694	2.136	1.824	1.271	2.167	2.191	1.351
1.6	0.826	0.835	1.282	2.967	0.857	2.35	2.82	2.194	3.038	2.234	1.369	3.056	3.008	1.475
1.7	0.765	0.795	1.418	4.035	0.803	3.00	3.60	2.744	4.137	2.590	1.440	4.083	3.789	1.553
1.8	0.708	0.744	1.565	5.382	0.769	3.72	4.16	3.331	5.440	2.863	1.490	5.250	4.510	1.604

\*Combinations of values of  $\sigma_g$  and  $n$  that are representative of atmospheric aerosols are indicated by an asterisk.

The above values were calculated at  $\lambda = 0.55\mu\text{m}$ . However, subsequent calculations have shown the values to be within 12% of those values for  $0.55\mu\text{m} \leq \lambda \leq 1.06\mu\text{m}$ .

TABLE 7. RANGE OF IMAGINARY INDEX,  $k$ , FOR WHICH THE VOLUME EXTINCTION COEFFICIENT IS CONSTANT TO WITHIN ABOUT 20% OR LESS

$r_g/\lambda^*$	Imaginary Index	Geometric Standard Deviation $\sigma_g$
0.01	0.001-0.0025 ( $\lambda = 10.6\mu\text{m}$ only)	2.0
0.01	0.001-0.01	2.0
0.01	0.001-0.02	2.5
0.05	0.001-0.005	1.5
0.05	0.001-0.025	2.0
0.05	0.001-0.1 ( $\lambda = 10.6\mu\text{m}$ only)	2.0
0.01	0.001-0.05	1.5
0.01	0.001-0.3	2.0
0.25	0.001-0.1	1.5
0.25	0.001-0.25	2.0
0.5	0.001-6.0 <sup>+</sup>	1.5
0.5	0.001-6.0 <sup>+</sup>	2.0
1.0	0.001-6.0 <sup>+</sup>	1.5
1.0	0.001-6.0 <sup>+</sup>	2.0

\*Applicable to wavelengths 0.55, 0.694, 1.06, 3.8 and 10.6 micrometers.  
<sup>+</sup>The upper limit of the imaginary index is only realistic at 10.6 $\mu\text{m}$  wavelength.

TABLE 8. VALUES OF IMAGINARY INDEX  $k_\lambda$  BELOW WHERE  
THE ABSORPTION COEFFICIENT VARIES LIN-  
EARLY (WITHIN 20%) WITH IMAGINARY INDEX  $k$   
WAVELENGTH =  $0.55\mu\text{m}$ ; REAL INDEX = 1.55

Geometric Mean Radius $r_g(\mu\text{m})$	Geometric Standard Deviation $\sigma_g$	$k_\lambda$
0.1	1.5	0.1
0.1	2.0	0.025
0.1	2.5	0.01
0.1	3.0	0.005
0.25	1.5	0.025
0.25	2.0	0.01
0.25	2.5	0.005
0.25	3.0	0.001
0.5	1.5	0.01
0.5	2.0	0.005
0.5	2.5	0.0025
1.0	1.5	0.005
1.0	2.0	0.0025
1.0	2.25	0.001

TABLE 9. VALUES OF IMAGINARY INDEX  $k_L$  BELOW WHERE  
THE ABSORPTION COEFFICIENT VARIES LINEARLY  
(WITHIN 20%) WITH IMAGINARY INDEX  $k$   
WAVELENGTH =  $1.06\mu\text{m}$ ; REAL INDEX = 1.6

Geometric Mean Radius $r_g(\mu\text{m})$	Geometric Standard Deviation $\sigma_g$	$k_L$
0.1	1.5	0.5
0.1	2.0	0.05
0.1	2.5	0.01
0.1	3.0	0.005
0.25	1.5	0.10
0.25	2.0	0.025
0.25	2.5	0.01
0.25	3.0	0.0025
0.5	1.5	0.025
0.5	2.0	0.01
0.5	2.5	0.005
1.0	1.5	0.005
1.0	2.0	0.005
1.0	2.25	0.001

TABLE 10. VALUES OF IMAGINARY INDEX  $k_c$ , ABOVE WHICH  
THE ABSORPTION COEFFICIENT IS CONSTANT (TO  
WITHIN ABOUT 40%) WITH IMAGINARY INDEX  
WAVELENGTH =  $0.55\mu\text{m}$ ; REAL INDEX = 1.55

Geometric Mean Radius $r_g$ ( $\mu\text{m}$ )	Geometric Standard Deviation $\sigma_g$	$k_c$
0.1	1.5	0.25
0.1	2.0	0.25
0.1	2.5	0.1
0.1	3.0	0.05
0.25	1.5	0.10
0.25	2.0	0.05
0.25	2.5	0.025
0.25	3.0	0.025
0.5	1.5	0.05
0.5	2.0	0.025
0.5	2.5	0.025
1.0	1.5	0.025
1.0	2.0	0.025
1.0	2.25	0.01

TABLE 11. VALUES OF IMAGINARY INDEX  $k_c$  ABOVE WHICH  
THE ABSORPTION COEFFICIENT IS CONSTANT (TO  
WITHIN ABOUT 40%) WITH IMAGINARY INDEX.  
WAVELENGTH =  $1.06\mu\text{m}$ ; REAL INDEX = 1.6

Geometric Mean Radius $r_g(\mu\text{m})$	Geometric Standard Deviation ( $\sigma_g$ )	$k_c$
0.1	1.5	1.0
0.1	2.0	0.5
0.1	2.5	0.1
0.1	3.0	0.05
0.25	1.5	0.25
0.25	2.0	0.10
0.25	2.5	0.05
0.25	3.0	0.025
0.5	1.5	0.1
0.5	2.0	0.05
0.5	2.5	0.025
1.0	1.5	0.05
1.0	2.0	0.025
1.0	2.25	0.025

TABLE 12. VALUES OF  $n_g^*$ ,  $\sigma_n^+$  WHICH SATISFY THE EQUATION

$$\sigma_E = \sigma_{E_{\max}} \exp \left\{ -\frac{1}{2} \left[ \frac{\ln(n/n_g)}{\ln \sigma_n} \right]^2 \right\} \text{ FOR A}$$

SELECTION OF  $r_g$ ,  $\sigma_g$ , AND  $k$  VALUES AT  $\lambda = 3.8\mu\text{m}$

$\lambda$	$r_g$	$\sigma_g$	$n_g$	$\sigma_n$	$\sigma_{E_{\max}}$	$k$
3.8	0.05	1.5	0.8	2.25	$6.34 \times 10^{-3}$	0.05
3.8	0.05	2.5	5.5	2.15	$9.0 \times 10^{-2}$	0.05
3.8	0.1	2.0	6.0	1.925	$1.1 \times 10^{-1}$	0.01
3.8	0.1	2.0	6.0	1.925	$1.1 \times 10^{-1}$	0.05
3.8	0.5	2.0	2.6	1.9	$5.2 \times 10^{-2}$	0.05
3.8	1.0	1.5	2.1	1.43	$6.64 \times 10^{-2}$	0.05

\* Peak extinction occurs at real index  $n = n_g$

+  $\sigma_n$  is the geometric standard deviation of the  $n$  values centered around  $n_g$ .

TABLE 13. RANGE OF IMAGINARY INDEX  $k$ , FOR WHICH THE EXTINCTION COEFFICIENT IS CONSTANT (TO WITHIN 15%). WAVELENGTH =  $3.8\mu\text{m}$ ; REAL INDEX  $n = 1.7$

Range of $k$ for Which Extinction is Constant	Geometric Mean Radius, $r_g$	Geometric Standard Deviation, $\sigma_g$
0.001 to 0.005	0.05	2.0
0.001 to 0.05	0.05	2.5
0.001 to 0.025	0.1	2.0
0.001 to 0.015	0.25	1.5
0.001 to 0.25	0.25	2.0
0.001 to 0.1	0.5	1.5
0.001 to 2.0	0.5	2.0
0.001 to 1.0	0.75	1.5
0.001 to 2.0	0.75	2.0
0.001 to 2.0	$r_g \geq 1.0$	$\sigma_g \geq 1.5$



TABLE 14. THE RATIO OF EXTINCTION COEFFICIENTS AT VALUES OF  
REAL INDEX  $n$ ,  $\sigma_E(n)/\sigma_E(n=1.4)$ , FOR TYPICAL  
AEROSOL SIZE DISTRIBUTIONS. WAVELENGTH  $\lambda = 3.8 \mu m$ ;  
IMAGINARY INDEX  $k = 0.025$

$n \backslash \sigma_g$		$\sigma_E(n)/\sigma_E(n=1.4)$											
		$r_g = 0.01 \mu m$			$r_g = 0.05 \mu m$			$r_g = 0.1 \mu m$			$r_g = 0.25 \mu m$		
		2.0	2.5	1.0	1.5	2.0	2.5	1.0	1.5	2.0	1.0	1.5	2.0
1.4	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0
1.6	0.86	0.93	0.87	1.12	1.88	0.96	1.68	1.64	1.55	1.80	1.30	2.0	1.76
1.8	0.74	0.86	0.77	1.29	2.93	0.94	2.69	2.14	2.28	2.53	1.48	3.3	2.37
2.0	0.63	0.85	0.67	1.49	3.96	0.94	3.99	2.54	3.18	3.08	1.59	4.65	2.78
2.25	0.52	0.83	0.57	1.77	5.03	0.95	5.79	2.93	4.72	3.60	1.66	6.13	3.09
2.5	0.43	0.82	0.49	2.07	5.84	0.98	7.67	3.19	7.02	3.97	1.69	7.13	3.24
												1.92	1.29
												1.64	1.27
												1.99	1.36
												2.09	1.36
												2.04	1.33
												1.92	1.29

TABLE 15. RANGE OF IMAGINARY INDEX  $k$  FOR WHICH  
EXTINCTION COEFFICIENT IS LINEAR (TO  
WITHIN 10%) WITH  $k$ ;  $n = 1.7$ ,  $\lambda = 3.8\mu\text{m}$

Range of $k$ , for Which $\sigma_E$ is Linear with $k$	$r_g$	$\sigma_g$
0.001 to 1.0	0.01	$\leq 2.0$
0.025 to 1.0	0.01	2.5
0.001 to 1.0	0.05	1.0*
0.01 to 1.0	0.05	1.5
0.05 to 1.0	0.05	2.0
0.01 to 1.0	0.1	1.0*
0.05 to 1.0	0.1	1.5
0.25 to 1.0	0.1	2.0
0.25 to 1.0	0.5	$\leq 1.125^*$

\*Values of  $\sigma_g$  do not represent realistic values for atmospheric aerosols.

TABLE 16. PARTICLE SIZE DISTRIBUTION PARAMETERS  
USED IN THE VOLUME EXTINCTION COEFFICIENT  
VERSUS REAL INDEX OF REFRACTION AT  $\lambda =$   
10.6 $\mu\text{m}$  (Figures 12-14).

Geometric Mean Radius $r_g$	Geometric Standard Deviation $\sigma_g$	Geometric Mean Radius $r_g$	Geometric Standard Deviation $\sigma_g$
0.005	1.0	0.5	1.0
0.005	1.5	0.5	2.0
0.005	2.0		
0.005	2.5	0.75	1.0
		0.75	1.5
0.01	1.0		
0.01	2.0	1.0	1.0
0.01	2.5	1.0	1.5
0.05	1.0	2.5	1.0
0.05	1.5	2.5	1.5
0.05	2.0		
0.05	2.5	5.0	1.0
		5.0	1.5
0.1	1.0		
0.1	2.0	10.0	1.0
		10.0	1.5
0.25	1.0		
0.25	1.5		
0.25	2.0		

TABLE 17. VALUES OF  $n_g$ ,  $\sigma_g$ , WHICH SATISFY THE EQUATION

$$\sigma_E = \sigma_{E_{\max}} \exp \left\{ -\frac{1}{2} \left[ \frac{\ln (n/n_g)}{\ln \sigma_n} \right]^2 \right\}$$

FOR A SELECTION OF  $r_g$ ,  $\sigma_g$  AND  $k$  VALUES AT  $\lambda = 10.6 \mu\text{m}$

$\lambda$	$r_g$	$\sigma_g$	$n_g$	$\sigma_{E_{\max}}$	$\sigma_n$	$k$
10.6	0.05	2.0	0.8	$4.38 \times 10^{-3}$	2.25	0.1
10.6	0.10	2.0	0.8	$4.33 \times 10^{-3}$	2.5	0.1
10.6	0.25	2.0	6.5	$3.73 \times 10^{-3}$	1.9	0.1
10.6	0.5	2.0	4.35	$3.7 \times 10^{-3}$	2.1	0.1
10.6	0.75	1.5	5.0	$5.27 \times 10^{-3}$	1.625	0.01
10.6	0.75	1.5	5.0	$5.27 \times 10^{-3}$	1.65	0.1
10.6	1.0	1.5	4.2	$5.6 \times 10^{-3}$	1.65	0.1
10.6	2.5	1.5	2.25	$2.62 \times 10^{-3}$	1.5	0.1

TABLE 18. VALUES OF THE RATIO OF EXTINCTION COEFFICIENTS  
 $\sigma_E(n)/\sigma_E(n=1.2)$  AT WAVELENGTH  $\lambda = 10.6 \mu\text{m}$   
 FOR TYPICAL AEROSOL SIZE DISTRIBUTIONS.  
 IMAGINARY INDEX  $k = 0.05$

$\frac{\sigma_g}{n}$	$\sigma_E(n)/\sigma_E(n=1.2)$																						
	$r_g = 0.1 \mu\text{m}$			$r_g = 0.25 \mu\text{m}$			$r_g = 0.5 \mu\text{m}$			$r_g = 0.75 \mu\text{m}$			$r_g = 1.0 \mu\text{m}$			$r_g = 2.5 \mu\text{m}$							
	1.5	2.0	2.5	1.5	2.0	2.5	1.5	2.0	2.5	1.5	2.0	2.5	1.5	2.0	2.5	1.5	2.0	2.5					
1.2	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0	1.0					
1.4	0.88	0.93	1.42		0.91	1.28	2.10		1.07	1.93	1.94		1.33	2.19	1.73		1.63	2.21	1.59		2.4	1.63	1.22
1.6	0.76	0.86	2.15		0.83	1.74	3.37		1.19	3.34	2.66		1.88	3.64	2.19		2.68	3.38	1.91		3.99	1.89	1.27
1.8	0.65	0.80	3.18		0.75	2.42	4.43		1.36	4.96	3.17		2.61	4.90	2.47		4.20	4.24	2.08		5.12	1.97	1.28
2.0	0.55	0.76	4.37		0.69	3.36	5.29		1.57	6.49	3.53		3.58	5.87	2.63		6.08	4.86	2.16		5.64	1.98	1.27
2.25	0.45	0.73	6.0		0.63	4.86	6.15		1.87	8.03	3.80		5.30	6.78	2.74		8.77	5.33	2.21		5.74	1.96	1.25
2.5	0.37	0.72	7.47		0.59	6.43	6.76		2.20	9.18	3.95		7.46	7.37	2.79		11.59	5.58	2.21		5.54	1.92	1.28
3.5	0.18	0.74	10.85		0.51	11.56	8.11		6.01	11.60	4.08		16.45	8.05	2.75		18.07	5.65	2.15		4.64	1.78	1.18

TABLE 19. RANGE OF IMAGINARY INDEX  $k$ , FOR WHICH THE VOLUME  
EXTINCTION COEFFICIENT IS CONSTANT (TO WITHIN 15%)  
FOR REAL INDEX  $n = 1.8$ , WAVELENGTH  $\lambda = 10.6\mu\text{m}$

Range of $k$ for Which Extinction is Constant	Geometric Mean Radius $r_g$	Geometric Standard Deviation $\sigma_g$
0.001 to 0.0025	0.1	2.0
0.001 to 0.035	0.1	2.5
0.001 to 0.025	0.25	2.0
0.001 to 0.25	0.25	2.5
0.001 to 0.01	0.5	1.5
0.001 to 0.1	0.5	2.0
0.001 to 1.0	0.5	2.5
0.001 to 0.025	0.75	1.5
0.001 to 0.5	0.75	2.0
0.001 to 8.0	0.75	2.5
0.001 to 0.05	1.0	1.5
0.001 to 1.0	1.0	2.0
0.001 to 8.0	1.0	2.5
0.001 to 1.0	2.5	$\sigma_g \leq 1.5$
0.001 to 8.0	$r_g > 2.5$	$\sigma_g > 1.0$

TABLE 20. RANGE OF IMAGINARY INDEX  $k$  FOR WHICH  
EXTINCTION COEFFICIENT IS LINEAR WITHIN  
15%, WITH  $k$ ; FOR  $n = 1.8$   $\lambda = 10.6\mu\text{m}$ .  
THE SIZE DISTRIBUTION REPRESENT REALIS-  
TIC VALUES FOR ATMOSPHERIC AEROSOLS

Range of $k$ , Over Which $\sigma_E$ is Linear with $k$	$r_g$ ( $\mu\text{m}$ )	$\sigma_g$
0.001 to 1.0	0.01	$\leq 2.5$
0.001 to 1.0	0.05	$\leq 1.5$
0.01 to 1.0	0.05	2.0
0.01 to 1.0	0.05	2.5
0.001 to 1.0	0.1	1.5
0.05 to 1.0	0.1	2.0
* * *	0.1	2.5
0.025 to 1.0	0.25	1.5
0.25 to 1.0	0.25	2.0
* * *	0.25	2.5
0.25 to 1.0	0.5	1.5
* * *	0.5	2.0
* * *	0.5	2.5
0.25 to 1.0	0.75	1.5
* * *	0.75	2.0
* * *	0.75	2.5
* * *	1.0	1.5
* * *	1.0	2.0
* * *	$r_g > 1.0$	$\sigma_g > 1.5$

\* \* \*Denotes the range of  $k$ ,  $0.001 \leq k \leq 8.0$  over which  $\sigma_E$  is not linear  
with  $k$ .

TABLE 21. THE RELATIVE SENSITIVITY OF THE EXTINCTION COEFFICIENT TO CHANGES IN REAL INDEX  $n$  AND IMAGINARY INDEX  $k$  AT WAVELENGTH  $\lambda = 10.6\mu\text{m}$

Geometric Mean Radius $r_g$ ( $\mu\text{m}$ )	Geometric Standard Deviation $\sigma_g$					
	Extinction Dependence on $k$ ( $n = 1.8$ )			Extinction at $n = 2.5$ Extinction at $n = 1.25$		
	1.5	2.0	2.5	1.5	2.0	2.5
0.1	linear for $k > 0.001$	linear for $k > 0.05$	constant (within 15%) for $k: 0.001-0.25$	0.38	0.48	1.93
0.25	linear for $k: 0.05-1.0$	constant (within 15%) for $k: 0.001-0.025$ linear over $k: 0.25-1.0$	constant for $k:$ $0.001-0.25$	0.44	1.73	2.36
0.5	constant (within 15%) for $k: 0.001-0.01$	constant for $k:$ $0.001-0.1$	constant for $k:$ $0.001-1.0$	0.90	2.64	2.04
0.75	constant (within 15%) for $k: 0.001-0.025$	constant for $k:$ $0.001-0.5$	constant for $k:$ $0.001-8.0$	2.05	2.57	1.77
1.0	constant (within 15%) for $k: 0.001-0.05$	constant for $k:$ $0.001-1.0$	constant for $k:$ $0.001-8.0$	2.95	2.33	1.59



TABLE 22. VALUES OF EXTINCTION COEFFICIENT  $\sigma_E$  AND ABSORPTION  $\sigma_A$  IN  $\text{km}^{-1}$  (FOR PARTICULATE MASS OF  $100\mu\text{g}/\text{m}^3$ )  
AT  $\lambda = 0.55\mu\text{m}$ ,  $1.06\mu\text{m}$ , AND  $10.6\mu\text{m}$ ;  $k = 0.01$ ,  $n = 1.7$  FOR A RANGE OF VALUES  $r_g$  AND  $\sigma_g$

$r = 0.01\mu\text{m}$								
$\sigma_g$	1.125*		2.0		2.5			
$\lambda(\mu\text{m})$	$\sigma_E$	$\sigma_A$	$\sigma_E$	$\sigma_A$	$\sigma_E$	$\sigma_A$	$\sigma_E$	$\sigma_A$
0.55	$5.835 \times 10^{-2}$	$5.60 \times 10^{-3}$	$3.526 \times 10^{-2}$	$6.425 \times 10^{-3}$	$2.265 \times 10^{-1}$	$1.10 \times 10^{-2}$		
1.06	$2.903 \times 10^{-3}$	$2.88 \times 10^{-3}$	$4.973 \times 10^{-3}$	$2.88 \times 10^{-3}$	$4.92 \times 10^{-2}$	$4.0 \times 10^{-3}$		
10.6	$2.87 \times 10^{-4}$	$2.87 \times 10^{-4}$	$2.71 \times 10^{-4}$	$2.71 \times 10^{-4}$	$2.64 \times 10^{-4}$	$2.59 \times 10^{-4}$		
$r = 0.05\mu\text{m}$								
	1.0*		1.125*		1.5		2.0	
0.55	$3.28 \times 10^{-2}$	$6.9 \times 10^{-3}$	$3.72 \times 10^{-2}$	$7.07 \times 10^{-3}$	$1.426 \times 10^{-1}$	$9.59 \times 10^{-3}$	$3.25 \times 10^{-1}$	$1.535 \times 10^{-2}$
1.06	$4.83 \times 10^{-3}$	$3.06 \times 10^{-3}$	$5.19 \times 10^{-3}$	$3.09 \times 10^{-3}$	$1.48 \times 10^{-2}$	$3.40 \times 10^{-3}$	$1.2 \times 10^{-1}$	$5.8 \times 10^{-3}$
10.6	$2.86 \times 10^{-4}$	$2.87 \times 10^{-4}$	$2.88 \times 10^{-4}$	$2.88 \times 10^{-4}$	$2.84 \times 10^{-4}$	$2.83 \times 10^{-4}$	$3.0 \times 10^{-4}$	$2.75 \times 10^{-4}$
$r = 0.1\mu\text{m}$								
	1.125*		1.5		2.0			
0.55	$2.16 \times 10^{-1}$	$1.13 \times 10^{-2}$	$4.0 \times 10^{-1}$	$1.56 \times 10^{-2}$	$2.45 \times 10^{-1}$	$1.67 \times 10^{-2}$		
1.06	$2.17 \times 10^{-2}$	$3.73 \times 10^{-3}$	$8.06 \times 10^{-2}$	$5.14 \times 10^{-3}$	$1.69 \times 10^{-1}$	$8.07 \times 10^{-3}$		
10.6	$2.90 \times 10^{-4}$	$2.89 \times 10^{-4}$	$2.93 \times 10^{-4}$	$2.85 \times 10^{-4}$	$4.97 \times 10^{-4}$	$2.88 \times 10^{-4}$		
$r = 0.5\mu\text{m}$								
	1.0*		1.125*		1.5			
0.55	$1.02 \times 10^{-1}$	$1.67 \times 10^{-2}$	$1.19 \times 10^{-1}$	$1.86 \times 10^{-2}$	$4.0 \times 10^{-2}$	$9.23 \times 10^{-3}$		
1.06	$2.36 \times 10^{-2}$	$9.8 \times 10^{-3}$	$2.35 \times 10^{-1}$	$9.9 \times 10^{-3}$	$4.56 \times 10^{-2}$	$6.62 \times 10^{-3}$		
10.6	$4.83 \times 10^{-4}$	$3.06 \times 10^{-4}$	$5.19 \times 10^{-4}$	$3.09 \times 10^{-4}$	$1.2 \times 10^{-2}$	$5.85 \times 10^{-4}$		
$r = 1.0\mu\text{m}$								
	1.125*		1.5					
0.55	$6.6 \times 10^{-2}$	$1.27 \times 10^{-2}$	$4.39 \times 10^{-2}$	$1.02 \times 10^{-2}$				
1.06	$6.18 \times 10^{-2}$	$9.53 \times 10^{-3}$	$4.66 \times 10^{-2}$	$7.53 \times 10^{-3}$				
10.5	$2.12 \times 10^{-3}$	$3.73 \times 10^{-4}$	$8.06 \times 10^{-3}$	$5.14 \times 10^{-4}$				
$r = 5.0\mu\text{m}$								
	1.125*		1.5					
0.55	$1.17 \times 10^{-2}$	$4.71 \times 10^{-3}$	$8.0 \times 10^{-3}$	$3.33 \times 10^{-3}$				
1.06	$1.21 \times 10^{-2}$	$3.92 \times 10^{-3}$	$8.19 \times 10^{-3}$	$2.94 \times 10^{-3}$				
10.6	$2.35 \times 10^{-2}$	$9.9 \times 10^{-4}$	$1.21 \times 10^{-2}$	$9.52 \times 10^{-4}$				
$r = 10.0\mu\text{m}$								
	1.125*		1.5					
0.55	$5.73 \times 10^{-3}$	$2.50 \times 10^{-3}$	$3.91 \times 10^{-3}$	$1.71 \times 10^{-3}$				
1.06	$5.85 \times 10^{-3}$	$2.37 \times 10^{-3}$	$3.98 \times 10^{-3}$	$1.67 \times 10^{-3}$				
10.6	$6.18 \times 10^{-3}$	$9.53 \times 10^{-3}$	$4.66 \times 10^{-4}$	$1.21 \times 10^{-3}$				

\*Appropriate for monodisperse or near monodisperse aerosol size distributions only.

TABLE 23. VARIATION OF EXTINCTION COEFFICIENT AND ABSORPTION COEFFICIENT  
 $\sigma_E, \sigma_A$  WITH WAVELENGTH IN THE FORM OF  $\sigma_E, \sigma_A \sim \lambda^y$ .

$r_g$	$\sigma_g$	Values of $y$					
		Wavelength Range ( $\mu m$ )					
		0.55 - 1.06		1.06 - 10.6		0.55 - 10.6	
		Extinction	Absorption	Extinction	Absorption	Extinction	Absorption
(i) $n = 1.7, k = 0.01$							
0.01	1.125	-1.06	-1.0	-1.0	-1.0	-	-1.26
0.01	2.0	-2.98	-1.22	-1.26	-1.03	-	-1.26
0.01	2.5	-2.32	-1.19	-2.27	-1.19	-	-1.26
0.1	1.125	-3.50	-1.69	No slope		No slope	
0.1	1.5	-2.44	-1.69				
0.1	2.0	-0.57	-1.10				
0.5	1.125	+1.0	-0.96	-	-1.51	No slope	
0.5	2.0	+0.2	-0.50	-	-	No slope	
1.0	1.125	-0.1	-0.44	No slope		No slope	
	1.5	+0.09	-0.46				
5.0	1.125	+0.05	-0.28	No slope		No slope	
	1.5	+0.03	-0.19				
10.0	1.125	+0.03	-0.07	0.08	-	0.03	
	1.5	0.0	0.0	-	-	-	
(ii) $n = 1.8, k = 0.1$							
0.01	1.125	-1.02	-1.01	-1.0	-1.0	-	-1.00
	2.0	-1.81	-1.24	-1.07	-1.03	-	-1.08
	2.5	-1.81	-1.37	-1.56	-1.19	-	-1.22
0.1	1.125	-2.79	-1.75	-1.33	-1.12	No slope	
	1.5	-1.87	-1.44	-	-	No slope	
	2.0	-0.42	-0.64	-1.80	-1.31	No slope	
1.0	1.125	+0.13	+0.12	No slope		No slope	
	1.5	+0.10	+0.14	No slope		No slope	
10.0	1.125	+0.03	+0.05	0.09	0.12	0.07	0.11
	1.5	+0.02	+0.04	0.07	0.11	0.06	0.09

TABLE 24. A COMPARISON OF  $\sigma_E$  AND  $\sigma_A$  AT  $\lambda = 1.06\mu m, 3.8\mu m$   
AND  $10.6\mu m$  FOR REALISTIC VALUES OF  $r_g, \sigma_g$  AND  $k$   
FOR ATMOSPHERIC AEROSOLS

$k$	$n$	$r_g$	$\sigma_g$	$\lambda$	$\sigma_E$	$\sigma_A$
0.01	1.7	0.5	2.0	1.06	$4.56 \times 10^{-2}$	$6.62 \times 10^{-3}$
0.025	1.8	0.5	2.0	3.8	$4.59 \times 10^{-2}$	$5.67 \times 10^{-3}$
0.05	1.8	0.5	2.0	10.6	$1.54 \times 10^{-2}$	$2.83 \times 10^{-3}$
0.005	1.7	0.5	2.0	1.06	$4.56 \times 10^{-2}$	$3.84 \times 10^{-3}$
0.01	1.8	0.5	2.0	3.8	$4.62 \times 10^{-2}$	$2.60 \times 10^{-3}$
0.025	1.8	0.5	2.0	10.6	$1.50 \times 10^{-2}$	$1.51 \times 10^{-3}$
0.01	1.7	0.5	2.0	1.06	$4.56 \times 10^{-2}$	$6.62 \times 10^{-3}$
0.05	1.8	0.5	2.0	3.8	$4.54 \times 10^{-2}$	$9.44 \times 10^{-3}$
0.25	1.8	0.5	2.0	10.6	$1.87 \times 10^{-2}$	$9.85 \times 10^{-3}$
0.005	1.7	1.0	1.5	1.06	$4.65 \times 10^{-2}$	$4.33 \times 10^{-3}$
0.01	1.8	1.0	1.5	3.8	$6.48 \times 10^{-2}$	$2.86 \times 10^{-3}$
0.025	1.8	1.0	1.5	10.6	$1.11 \times 10^{-2}$	$1.32 \times 10^{-3}$
0.01	1.7	1.0	1.5	1.06	$4.66 \times 10^{-2}$	$7.54 \times 10^{-3}$
0.025	1.8	1.0	1.5	3.8	$6.39 \times 10^{-2}$	$6.44 \times 10^{-3}$
0.05	1.8	1.0	1.5	10.6	$1.19 \times 10^{-2}$	$2.56 \times 10^{-3}$
0.01	1.7	1.0	1.5	1.06	$4.66 \times 10^{-2}$	$7.54 \times 10^{-3}$
0.01	1.8	1.0	1.5	3.8	$6.26 \times 10^{-2}$	$1.11 \times 10^{-2}$
0.25	1.8	1.0	1.5	10.6	$1.83 \times 10^{-2}$	$1.04 \times 10^{-2}$
0.005	1.7	0.25	2.0	1.06	$1.01 \times 10^{-1}$	$4.65 \times 10^{-3}$
0.01	1.8	0.25	2.0	3.8	$5.34 \times 10^{-2}$	$2.13 \times 10^{-3}$
0.025	1.8	0.25	2.0	10.6	$4.78 \times 10^{-3}$	$9.27 \times 10^{-4}$
0.01	1.7	0.25	2.0	1.06	$1.00 \times 10^{-1}$	$8.33 \times 10^{-3}$
0.025	1.8	0.25	2.0	3.8	$5.11 \times 10^{-2}$	$5.00 \times 10^{-3}$
0.05	1.8	0.25	2.0	10.6	$5.60 \times 10^{-3}$	$1.83 \times 10^{-3}$
0.01	1.7	0.25	2.0	1.06	$1.00 \times 10^{-1}$	$8.33 \times 10^{-3}$
0.05	1.8	0.25	2.0	3.8	$5.14 \times 10^{-2}$	$9.06 \times 10^{-3}$
0.25	1.8	0.25	2.0	10.6	$1.18 \times 10^{-2}$	$8.36 \times 10^{-3}$

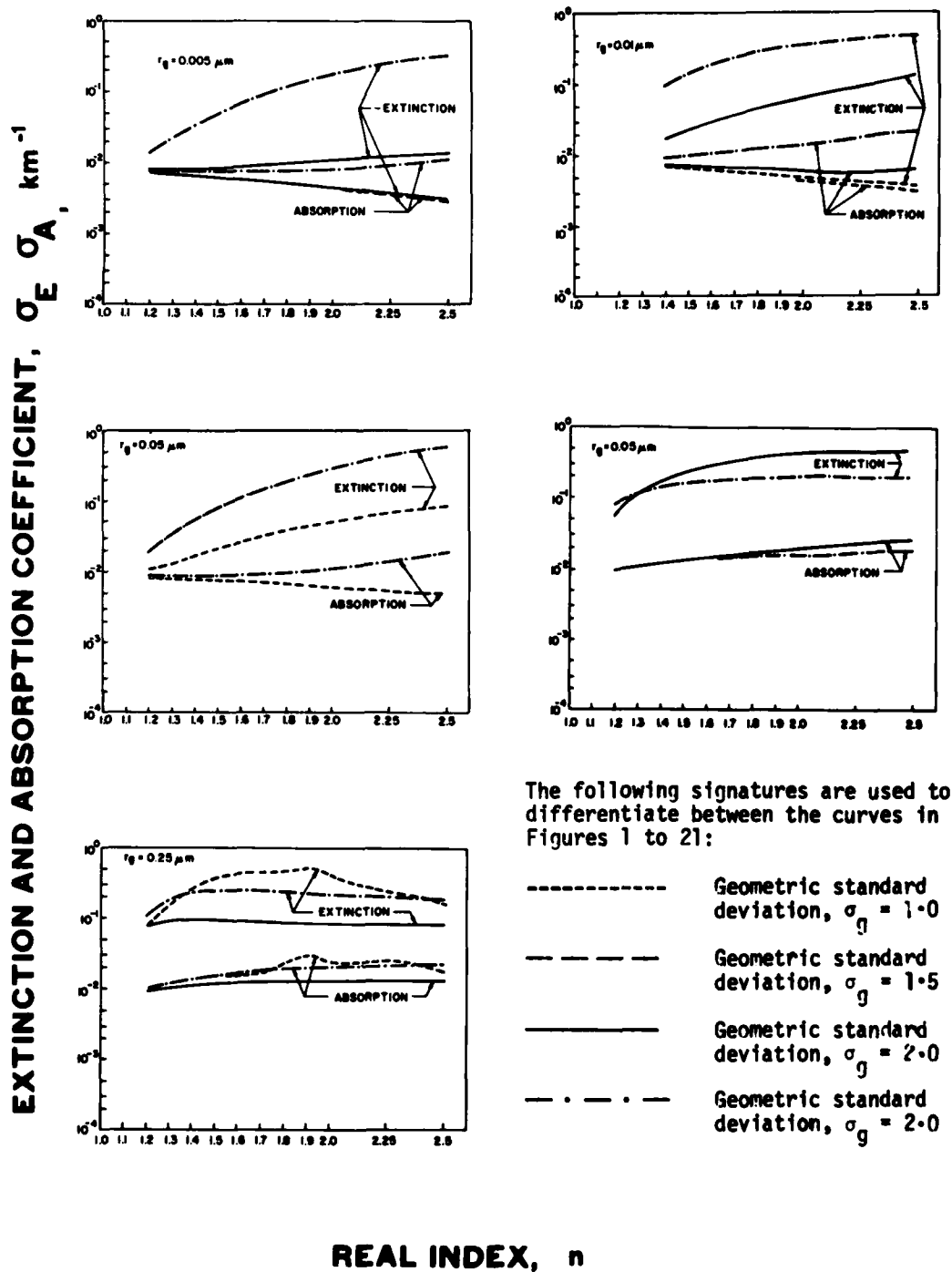


Figure 1. Volume extinction and absorption coefficients,  $\sigma_E$  and  $\sigma_A$   $\text{km}^{-1}$  as a function of real index  $n$ . Wavelength  $\lambda = 0.55 \mu\text{m}$ ; imaginary index  $k = 0.01$ .

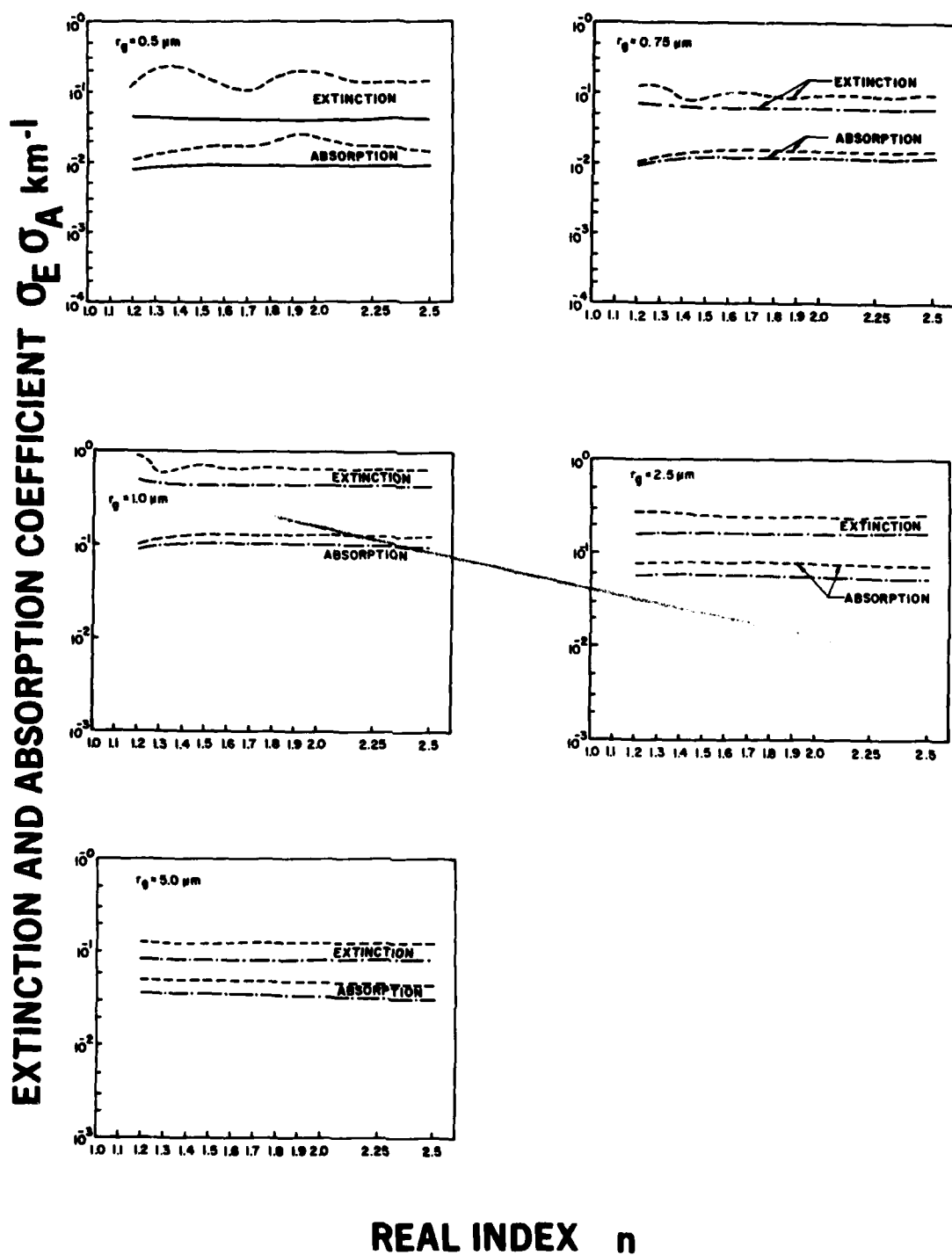


Figure 2. Volume extinction and absorption coefficients,  $\sigma_E$  and  $\sigma_A \text{ km}^{-1}$  as a function of real index  $n$ . Wavelength  $\lambda = 0.55 \mu\text{m}$ ; imaginary index  $k = 0.01$ .

EXTINCTION AND ABSORPTION COEFFICIENT,  $\sigma_E, \sigma_A \text{ km}^{-1}$

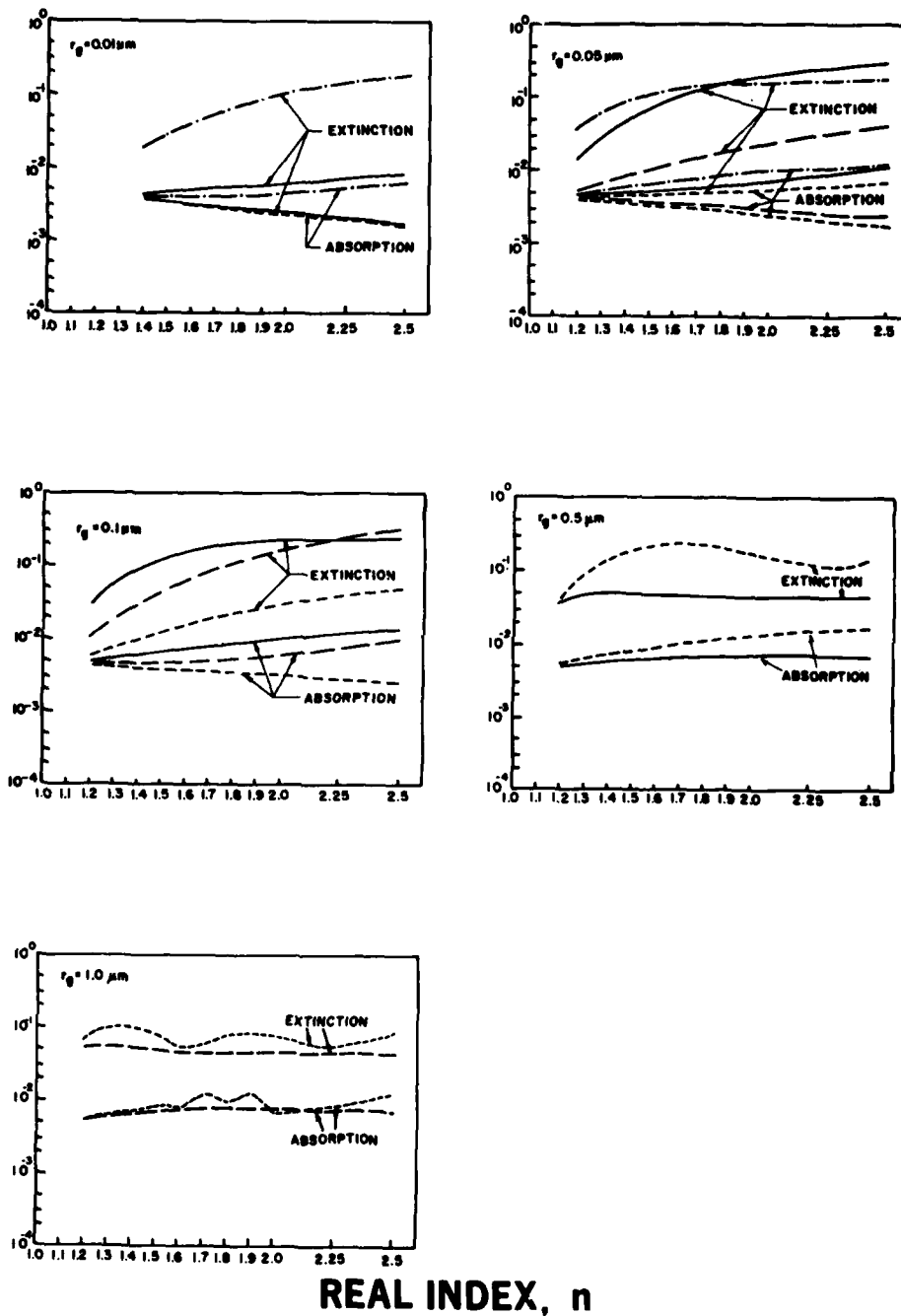


Figure 3. Volume extinction and absorption coefficients,  $\sigma_E$  and  $\sigma_A \text{ km}^{-1}$  as a function of real index  $n$ . Wavelength  $\lambda = 1.06 \mu\text{m}$ ; imaginary index  $k = 0.01$ .

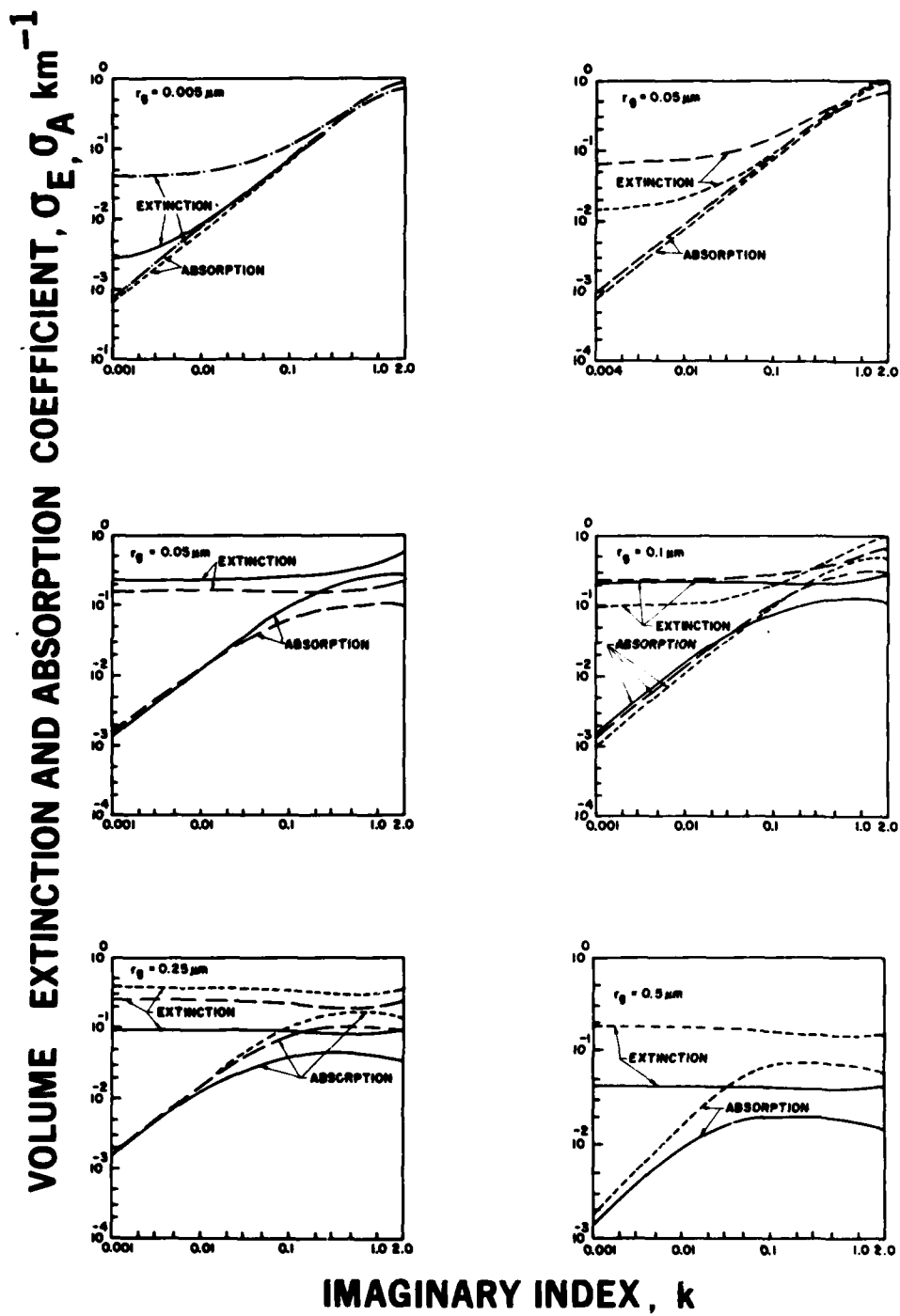
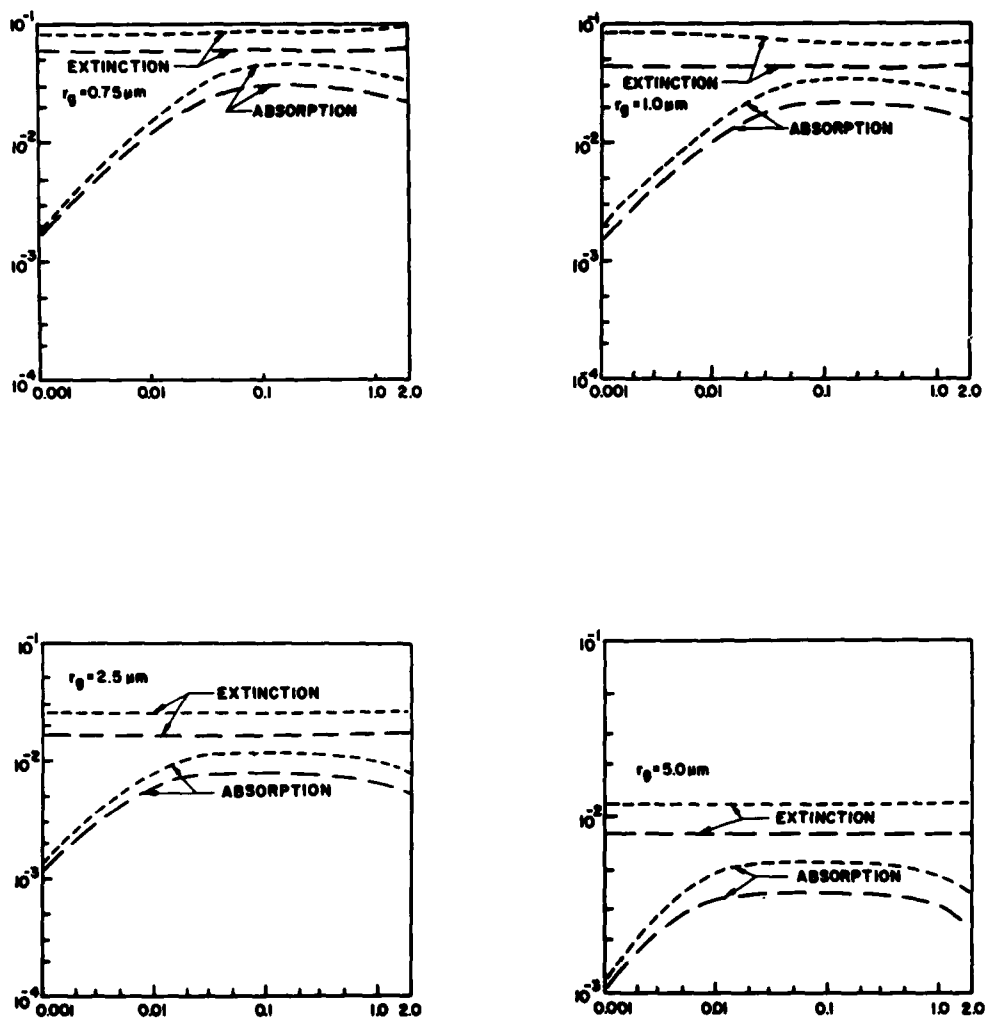


Figure 4. Volume extinction and absorption coefficients,  $\text{km}^{-1}$  as a function of imaginary index  $k$ . Wavelength  $\lambda = 0.55 \mu\text{m}$ ; real index  $n = 1.5$ .

# EXTINCTION AND ABSORPTION COEFFICIENT $\sigma_E, \sigma_A \text{ km}^{-1}$



## IMAGINARY INDEX, $k$

Figure 5. Volume extinction and absorption coefficients,  $\text{km}^{-1}$  as a function of imaginary index  $k$ . Wavelength  $\lambda = 0.55 \mu\text{m}$ , real index = 1.5.



# EXTINCTION AND ABSORPTION COEFFICIENT $\sigma_E, \sigma_A \text{ km}^{-1}$

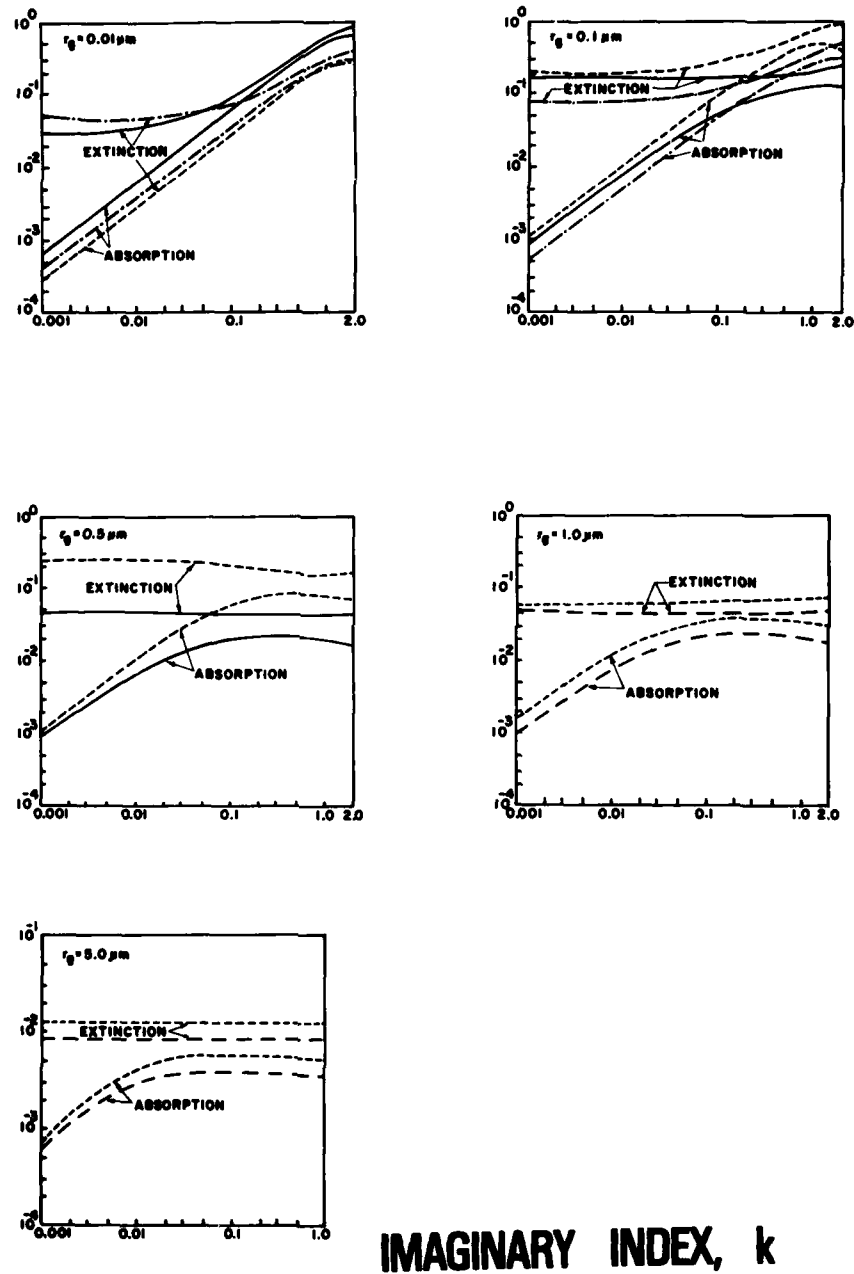


Figure 6. Volume extinction and absorption coefficients,  $\sigma_E$  and  $\sigma_A \text{ km}^{-1}$  as a function of imaginary index  $k$ . Wavelength  $\lambda = 1.06 \mu\text{m}$ ; real index = 1.7.

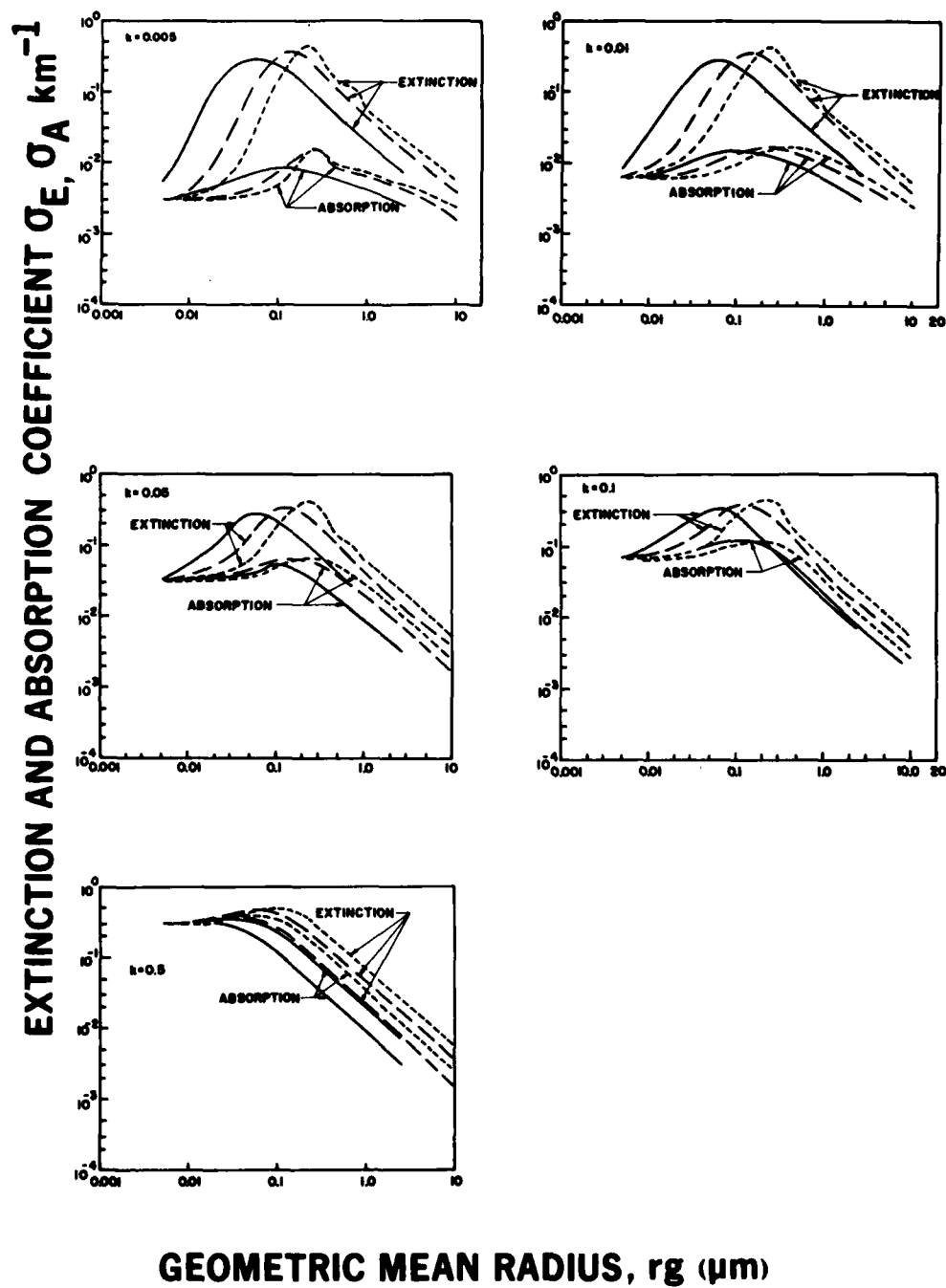
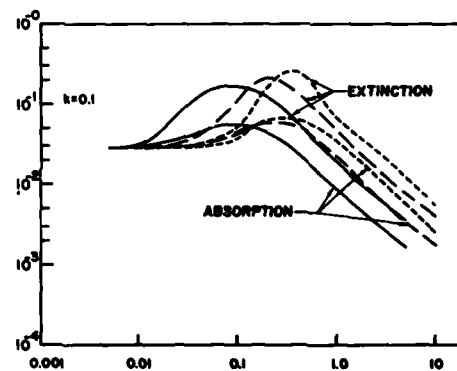
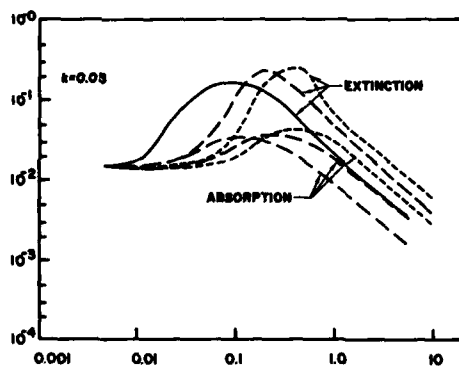
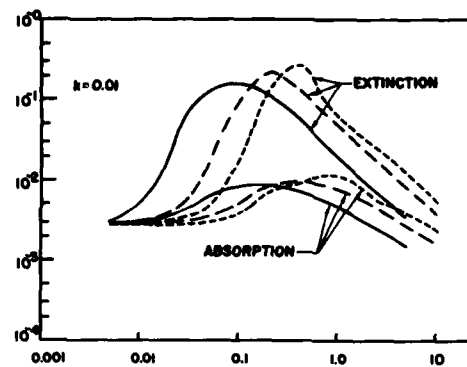
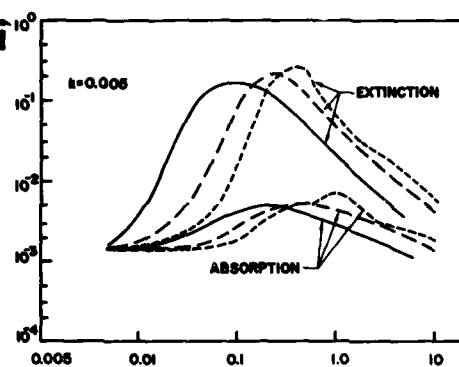


Figure 7. Extinction and absorption coefficient as a function of geometric mean radius  $r_g$ . Wavelength  $\lambda = 0.55\mu\text{m}$ ; real index = 1.6.

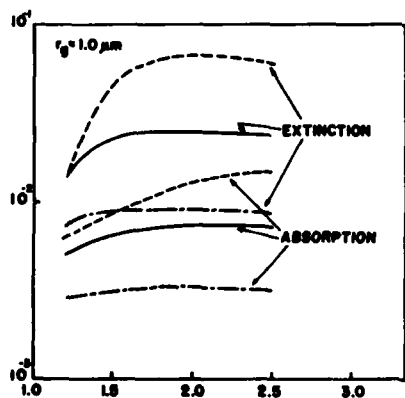
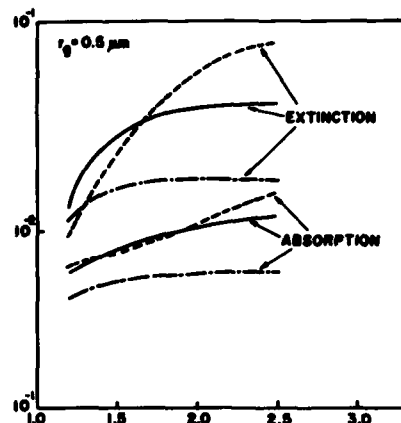
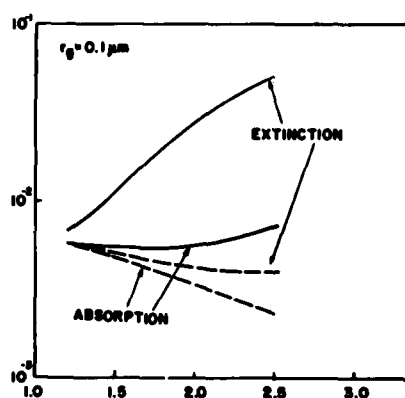
EXTINCTION AND ABSORPTION COEFFICIENT  $\sigma_E, \sigma_A \text{ km}^{-1}$



### GEOMETRIC MEAN RADIUS, $r_g(\mu\text{m})$

Figure 8. Volume extinction and absorption coefficient as a function of geometric mean radius  $r_g$ . Wavelength  $\lambda = 1.06\mu\text{m}$ ; real index = 1.7.

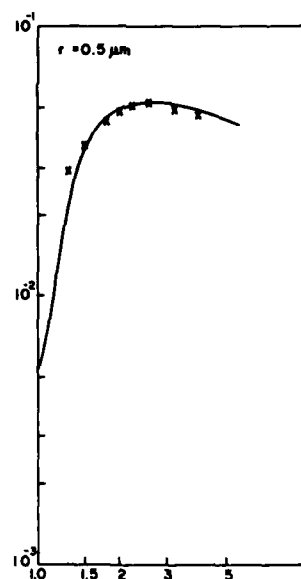
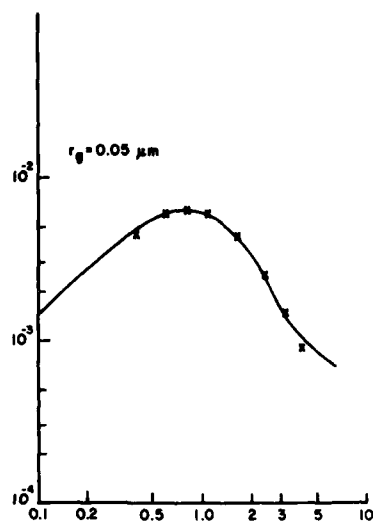
VOLUME EXTINCTION AND ABSORPTION COEFFICIENT  $\sigma_E, \sigma_A \text{ km}^{-1}$



REAL INDEX,  $n$

Figure 9. Volume extinction and absorption coefficient as a function of real index  $n$ . Wavelength  $\lambda = 3.8 \mu\text{m}$ ; imaginary index = 0.05.

VOLUME EXTINCTION COEFFICIENT  $\sigma_E \text{ km}^{-1}$



REAL INDEX,  $n$

Figure 10. Volume extinction coefficient as a function of real index. Wavelength  $\lambda = 3.8 \mu\text{m}$ ; imaginary index  $k = 0.05$ . Points marked x satisfy Eq. (20). For upper curve  $n_g = 0.8$ ,  $\sigma_n = 2.25$ ; for lower curve  $\sigma_g = 2.0$ ,  $n_g = 2.6$ ,  $\sigma_n = 1.9$ .

EXTINCTION AND ABSORPTION COEFFICIENT  $\sigma_E, \sigma_A \text{ km}^{-1}$

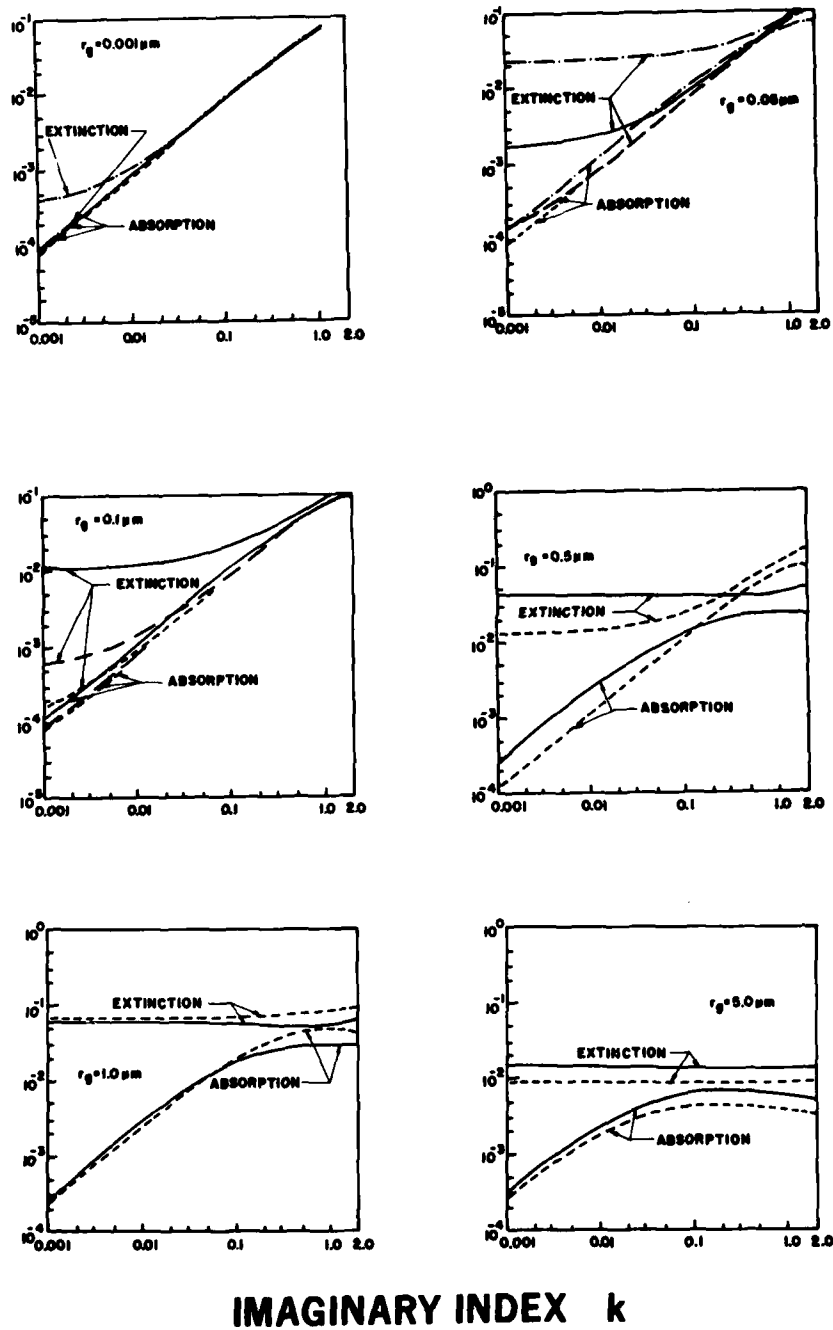


Figure 11. Volume extinction and absorption coefficient as a function of imaginary index  $k$ . Wavelength  $\lambda = 3.8 \mu\text{m}$ ; real index  $n = 1.7$ .

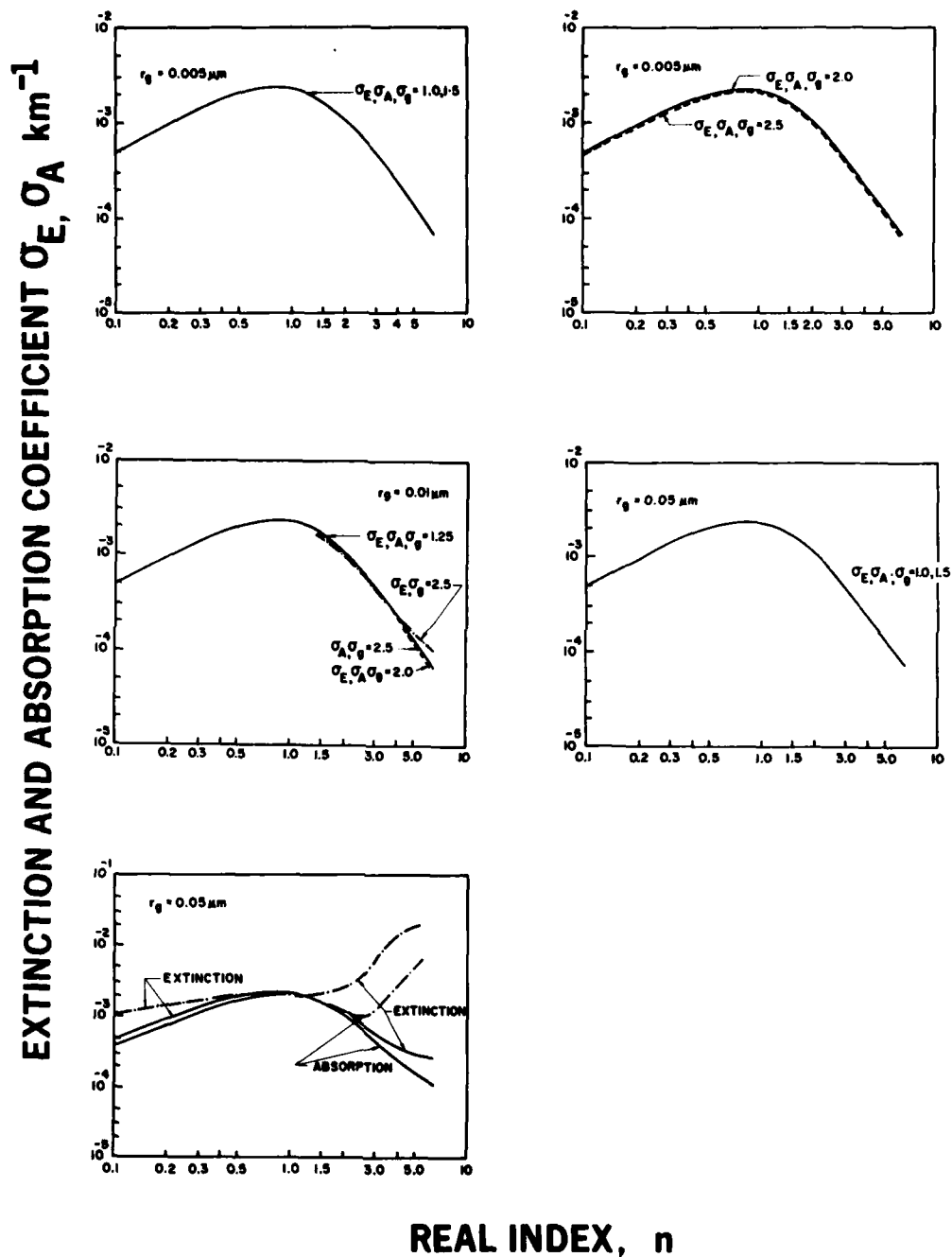


Figure 12. Volume extinction and absorption coefficient as a function of real index. Wavelength  $\lambda = 10.6 \mu\text{m}$ ; imaginary index  $k = 0.05$ .

# EXTINCTION AND ABSORPTION COEFFICIENT $\sigma_E, \sigma_A \text{ km}^{-1}$

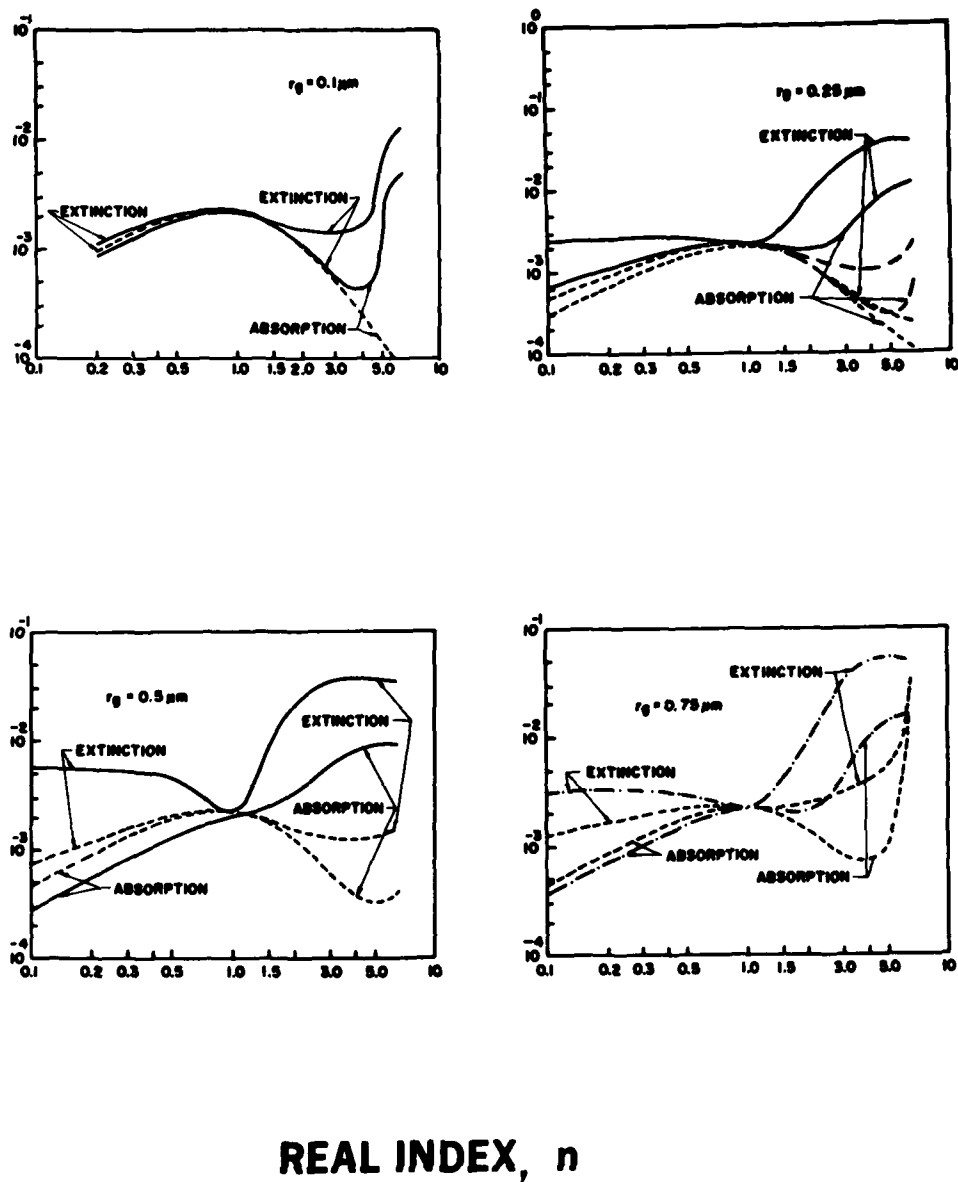
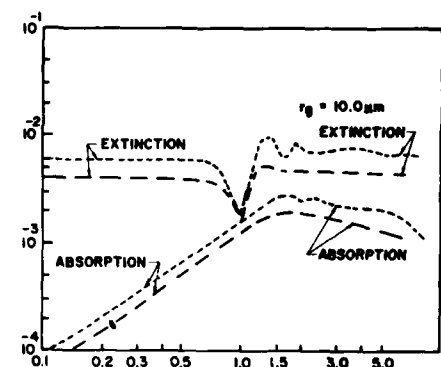
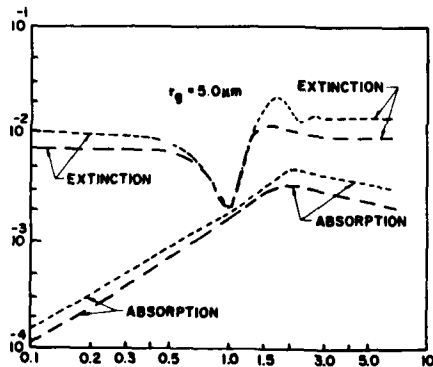
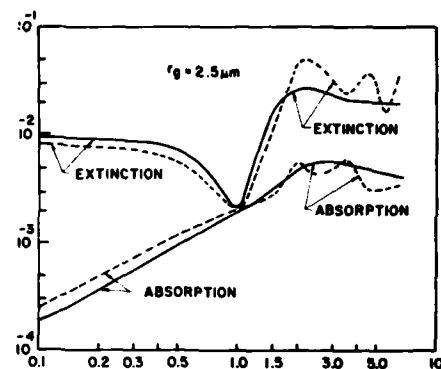
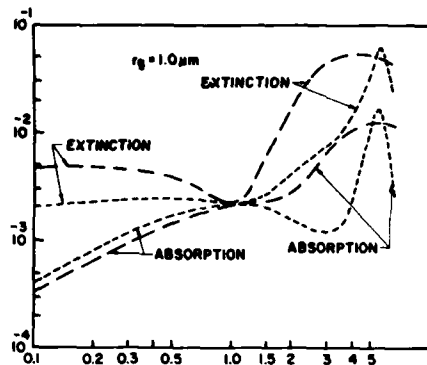


Figure 13. Volume extinction and absorption coefficient as a function of real index. Wavelength  $\lambda = 10.6 \mu\text{m}$ ; imaginary index = 0.05.



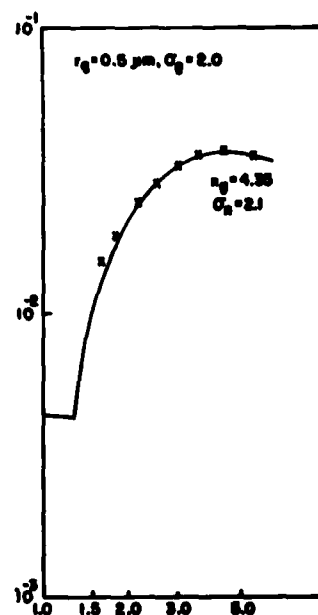
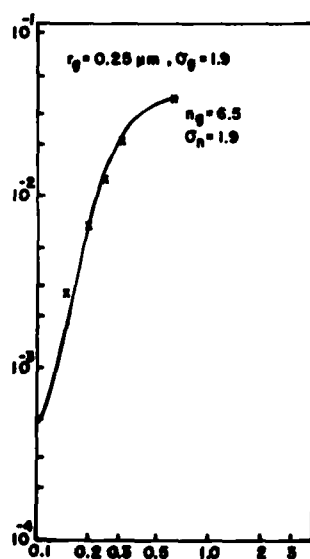
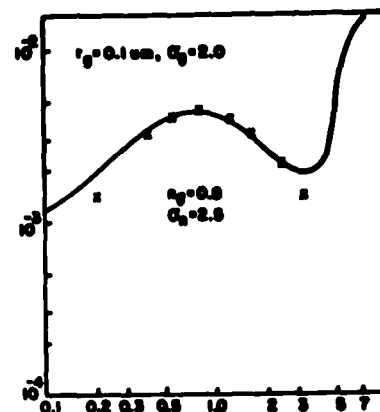
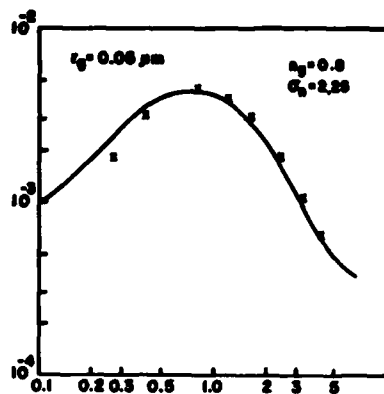
EXTINCTION AND ABSORPTION COEFFICIENT  $\sigma_E, \sigma_A$  km<sup>-1</sup>



REAL INDEX,  $n$

Figure 14. Volume extinction and absorption coefficient as a function of real index. Wavelength  $\lambda = 10.6 \mu\text{m}$ ; imaginary index  $k = 0.05$ .

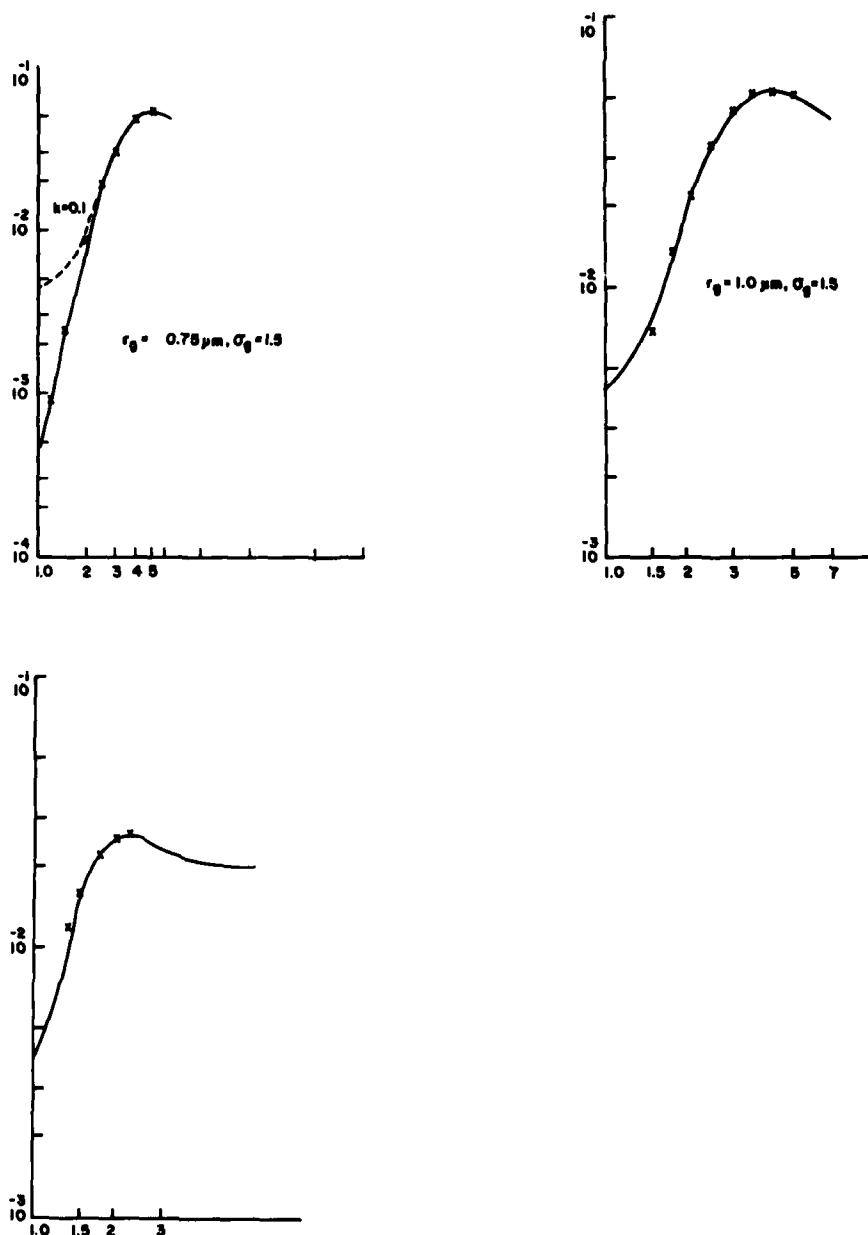
EXTINCTION COEFFICIENT,  $\sigma_E \text{ Km}^{-1}$



REAL INDEX  $n$

Figure 15. Volume extinction coefficient as a function of real index. Wavelength  $\lambda = 10.6 \mu\text{m}$ ; imaginary index  $k = 0.1$ . Points marked  $x$  satisfy Eq. (20).

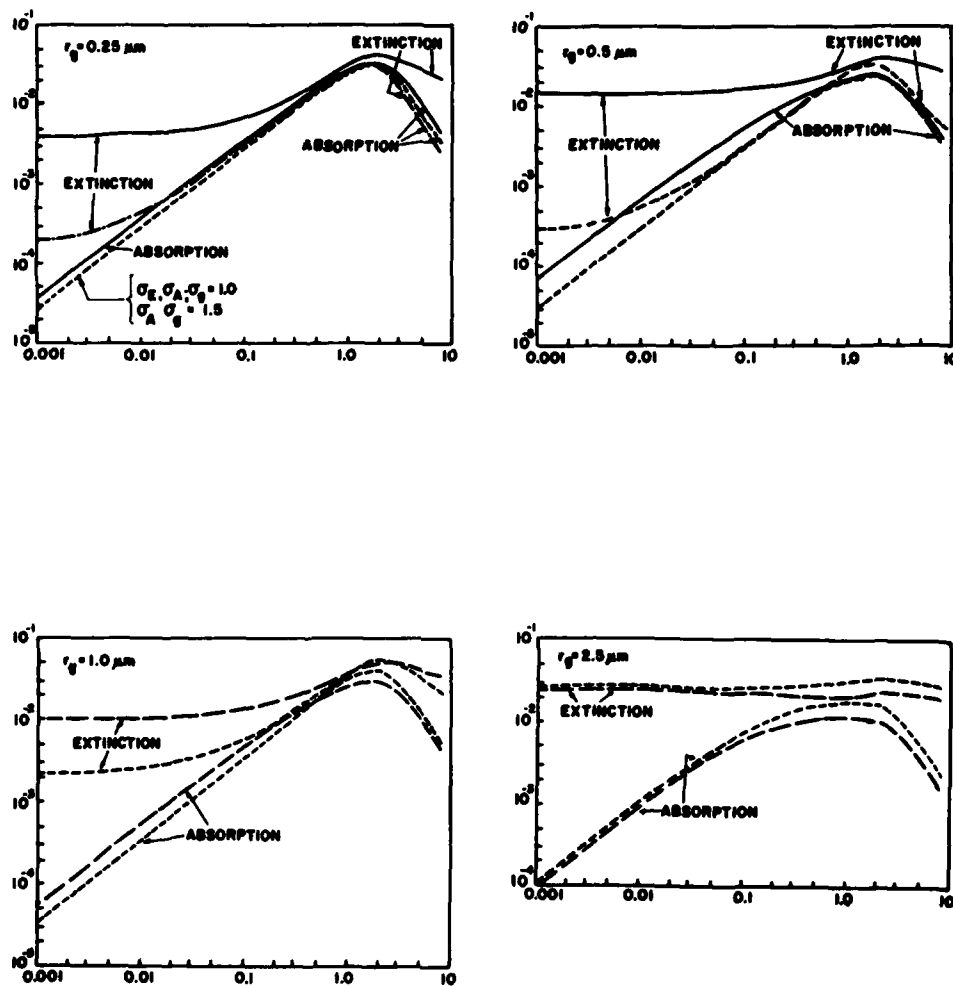
VOLUME EXTINCTION COEFFICIENT  $\sigma_E \text{ km}^{-1}$



REAL INDEX,  $n$

Figure 16. Volume extinction coefficient as a function of real index. Wavelength =  $10.6 \mu\text{m}$ ; imaginary index = 0.1. Points marked x satisfy Eq. (20).  
 Top left curve:  $n_g = 5.0, \sigma_n = 1.625$  for  $k = 0.01, \sigma_n = 1.65$  for  $k = 0.1$  (dashed curve).  
 Top right curve:  $n_g = 4.2, \sigma_n = 1.65$   
 Bottom curve:  $r_g = 2.5 \mu\text{m}, \sigma_g = 1.5; n_g = 2.25, \sigma_n = 1.5$

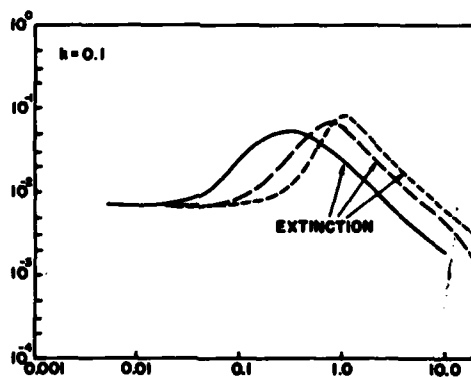
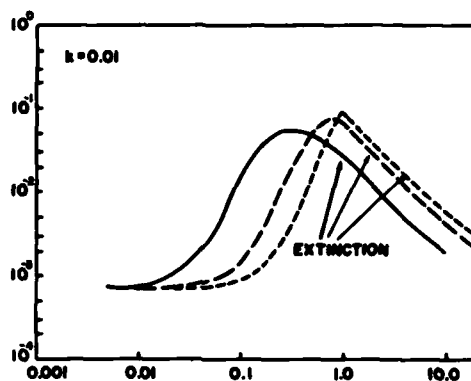
EXTINCTION AND ABSORPTION COEFFICIENT  $\sigma_E \sigma_A \text{ km}^{-1}$



IMAGINARY INDEX  $k$

Figure 17. Volume extinction and absorption coefficient as a function of imaginary index. Wavelength =  $10.6 \mu\text{m}$ ; real index = 1.8.

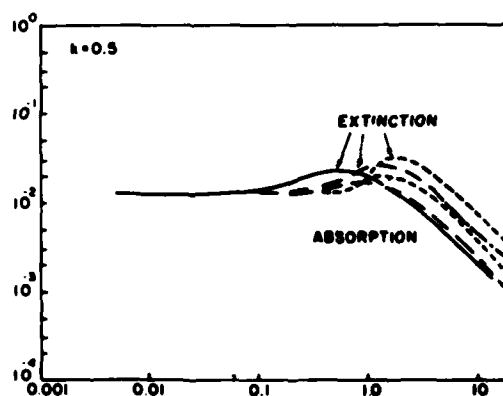
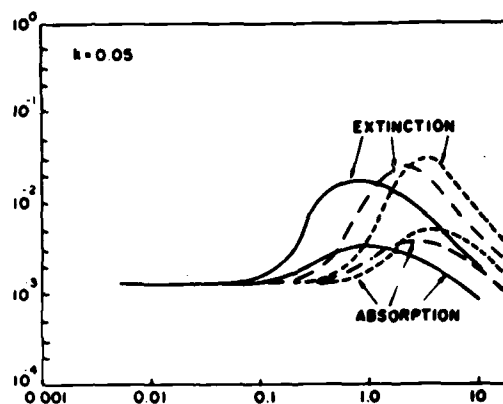
VOLUME EXTINCTION AND ABSORPTION COEFFICIENT  $\sigma_E \sigma_A \text{ km}^{-1}$



GEOMETRIC MEAN RADIUS,  $r_g$  ( $\mu\text{m}$ )

Figure 18. Volume extinction coefficient as a function of geometric mean radius. Wavelength =  $3.8 \mu\text{m}$ ; real index = 1.8.

EXTINCTION AND ABSORPTION COEFFICIENT,  $\sigma_E, \sigma_A \text{ km}^{-1}$



GEOMETRIC MEAN RADIUS,  $r_g (\mu\text{m})$

Figure 19. Volume extinction and absorption coefficient as a function of geometric mean radius. Wavelength =  $10.6 \mu\text{m}$ ; real index = 1.8.

VOLUME EXTINCTION AND ABSORPTION COEFFICIENT,  $\sigma_E, \sigma_A \text{ km}^{-1}$

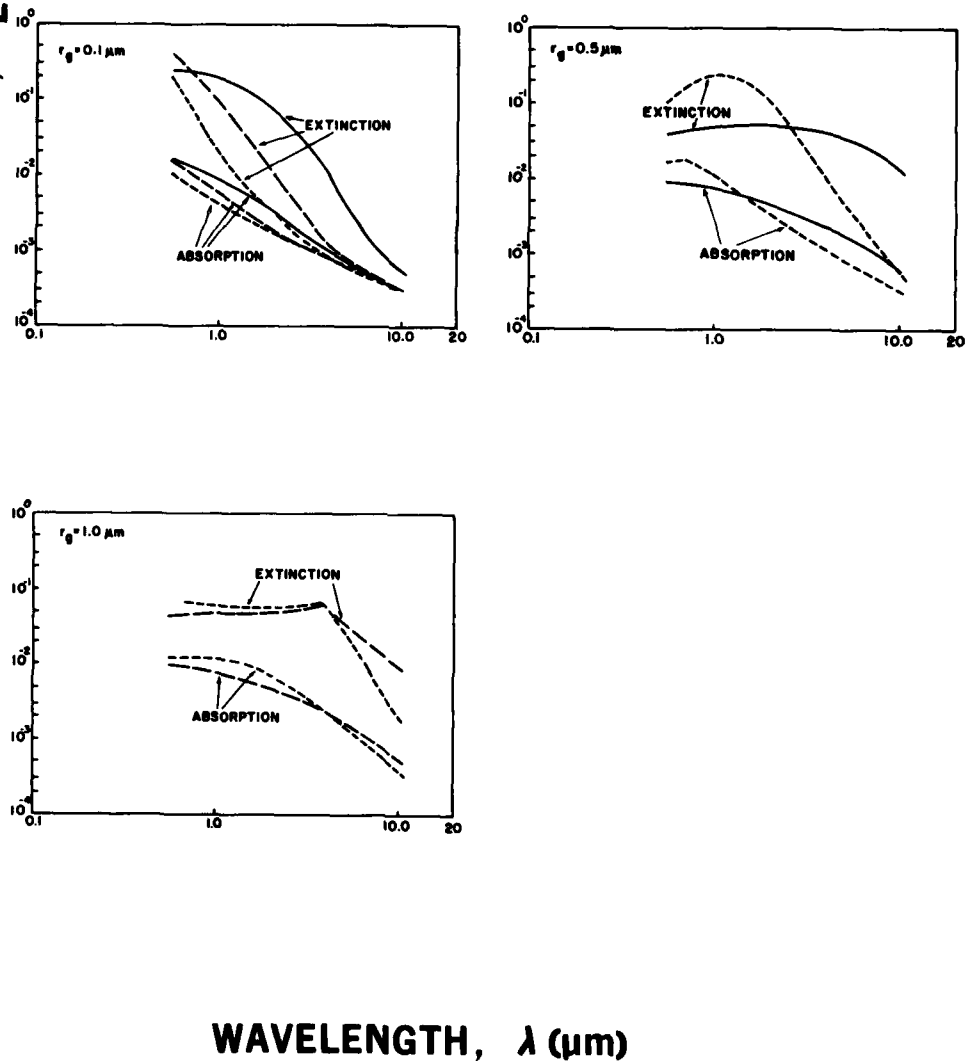
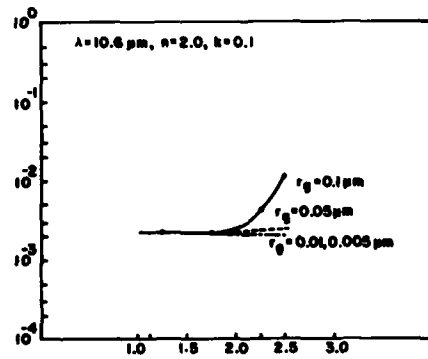
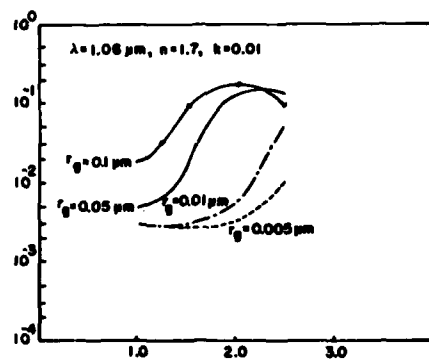
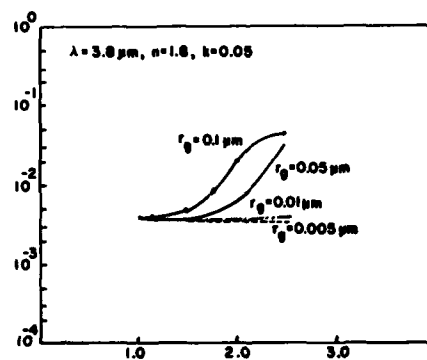
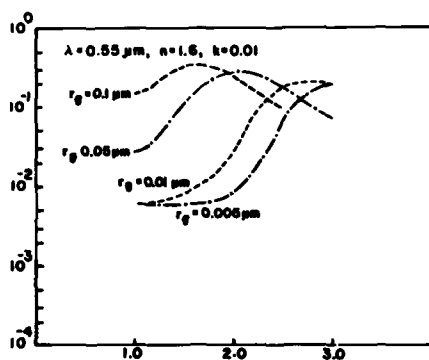


Figure 20. The variation of extinction and absorption coefficient with wavelength. Real index = 1.7, imaginary index = 0.01.

EXTINCTION AND ABSORPTION COEFFICIENT  $\sigma_E, \sigma_A \text{ km}^{-1}$



GEOMETRIC STANDARD DEVIATION,  $\sigma_g$

Figure 21. The variation of volume extinction coefficient with geometric standard deviation.



## REFERENCES

1. Mie, G., 1908, "Beitraeg Zur Optik Trueber Medien," Ann. Physik., 25:377-445.
2. van de Hulst, H. C., 1957, Light Scattering by Small Particles, New York, Wiley.
3. Deirmendjian, D., 1969, Electromagnetic Scattering on Spherical Polydispersions, New York, Elsevier.
4. Kerker, M., 1969, The Scattering of Light and Other Electromagnetic Radiation, New York, Academic Press.
5. Blättner, W. G. M., 1977, Research Note RRA-N-7703, Radiation Research Associates.
6. Grams, G. W., 1978, private communication.
7. Patterson, E. M., and D. A. Gillette, 1977, "Commonalities in Measured Size Distributions for Aerosols having a Soil-Derived Component," J. Geophys. Res., 82:2074-2082.
8. Whitby, K. T., W. E. Clark, V. A. Marple, G. M. Sverdrup, G. J. Sem, K. Willeke, B. Y. H. Liu, and D. Y. H. Pui, 1975, "Characterization of California Aerosols - I: Size Distributions of Freeway Aerosols," Atmospheric Environment, 9:463-482.
9. Davies, C. N., 1974, "Size Distributions of Atmospheric Particles," Aerosol Science, 5:293-300.
10. Hanel, G., 1968, "The Real Part of the Mean Complex Refractive Index and the Mean Density of Samples of Atmospheric Aerosol Particles," Tellus, 20:371-379.
11. Chylek, Peter, 1973, "Large-Sphere Limits of Mie-Scattering Functions," J. Opt. Soc. Am., 63:699-706.
12. Toon, O. B., J. B. Pollack, and B. N. Khare, 1976, "Optical Constants of Several Atmospheric Aerosol Species, Ammonium Sulphate, Aluminum Oxide and Sodium Chloride," J. Geophys. Res., 81:5733-5748.
13. Downing, H. D., L. W. Pinkley, P. P. Sethna, and D. Williams, 1974, "Optical Constants of Ammonium Sulphate in the Infrared," J. Opt. Soc. Am., 67:186-190.

14. Querry, M. R., W. E. Holland, and R. C. Waring, 1976, "Complex Refractive Index in the Infrared for NaCl, Na NO<sub>3</sub> and Na HCO<sub>3</sub> in Water," J. Opt. Soc. Am., 66:830-836.
15. Hale, G. M., and M. R. Querry, 1973, "Optical Constants of Water in the 200-nm to 200- $\mu$ m Wavelength Region," Applied Optics, 12:555-563.
16. Spitzer, W. G., and D. A. Kleinman, 1961, "Infrared Lattice Bands of Quartz," Phys. Rev., 121:1324.
17. Peterson, J. T., and J. A. Weinman, 1969, "Optical Properties of Quartz Dust Particles at Infrared Wavelengths," J. Geophys. Res., 74:6947-6952.
18. Palmer, K. F., and D. Williams, 1975, "Optical Constants of Sulphuric Acid; Application to the Clouds of Venus," Applied Optics, 14:208-219.
19. Foster, P. J. and C. R. Howarth, 1968, "Optical Constants of Carbons and Coals in the Infrared," Carbon, 6, pp. 719-729.
20. McCartney, J. T., J. B. Yasinsky, and S. Ergun, 1965, "Optical Constants of Coals," Fuel, 44, pp. 349-354.
21. Twitty, J. T. and J. A. Weinman, 1971, "Radiative Properties of Carbonaceous Aerosols," J. Appl. Meteorol., 10:725-731.
22. Dalzell, W. H. and A. F. Sarofim, 1969, "Optical Constants of Soot and their Application to Heat Flux Calculations," J. Heat Transfer 91:100-104.
23. Tyler, I. L., and M. R. Querry, 1977, "Complex Refractive Index in the Infrared for ZnCl<sub>2</sub> in Water," Phys. Rev.
24. Querry, M. R., G. Osborne, K. Lies, R. Jurdan, and R. M. Coveney, 1977, "Complex Refractive Index of Limestone in the Visible and Infrared," Applied Optics.
25. Toon, O. B., J. B. Pollack, and C. Sagan, 1977, "Physical Properties of the Particles Composing the Martian Dust Storm of 1971-1972," Icarus, 30:663-696.
26. Fischer, K., 1973, "Mass Absorption Coefficient of Natural Aerosol Particles in the 0.4 $\mu$ m-2.4 $\mu$ m Wavelength Interval," Beit. Phys. Atmos., 46:89-100.
27. Fischer, K., 1975, "Mass Absorption Indices of Various Types of Natural Aerosol Particles in the Infrared," Applied Optics, 14:2851-2856.

28. Lindberg, J. D., and L. S. Laude, 1974, "Measurement of the Absorption Coefficient of Atmospheric Dust," Applied Optics, 13:1923-1927.
29. Herman, B. M., R. S. Browning, J. J. DeLuisi, 1975, "Determination of the Effective Imaginary Term of the Complex Index of Atmospheric Dust by Remote Sensing: The Diffuse-Direct Radiation Method," J. Atmospheric Sci., 32:918-925.
30. Grams, G. W., I. H. Blifford, D. A. Gillette, and P. B. Russel, 1974, "Complex Index of Refraction of Airborne Soil Particles," J. Appl. Meteorol., 13:459-471.
31. Patterson, E. M., D. A. Gillette, and B. H. Sfocklin, 1977, "Complex Index of Refraction between 300 and 700 nm for Saharan Aerosols," Applied Optics (in press).
32. Hoidale, G. B., and A. J. Blanco, 1969, "Infrared Absorption Spectro of Atmospheric Dust over an Interior Desert Basin," Pure and Applied Geophys., 74:151-164.
33. Sverdrup, G. M., K. T. Whitby, and W. E. Clark, 1975, "Characterization of California Aerosols - II: Aerosol Size Distribution Measurements in the Mojave Desert," Atmospheric Environment, 9:483-494.
34. Hoihjelle, D. L., R. G. Pinnick, J. D. Lindberg, R. B. Loveland, E. B. Stenmark, C. J. Petracca, 1976, "Balloon-Borne Aerosol Particle Counter Measurements Made in Wintertime at Grafenwöhr, West Germany," ECOM-DR-76-3, Atmospheric Sciences Laboratory, US Army Electronics Command, White Sands Missile Range, NM.
35. De Luisi, J. J., P. M. Furukana, D. A. Gillette, B. G. Schuster, R. T. Charlson, W. M. Porch, R. W. Fegley, B. M. Herman, R. A. Rabinoff, J. T. Twitty, and J. A. Weinman, 1976, "Results of a Comprehensive Atmospheric Aerosol-Radiation Experiment in the Southwestern United States, Part I: Size Distribution, Extinction Optical Depth and Vertical Profiles of Aerosols Suspended in the Atmosphere," J. Appl. Meteorol., 15:441-454.
36. Meszaros, A., and K. Vissy, 1974, "Concentration, Size Distribution and Chemical Nature of Atmospheric Aerosol Particle in Remote Oceanic Areas," Aerosol Science, 5:101-109.
37. Jennings, S. G., 1977, "Mie Theory Sensitivity Studies - The Effects of Aerosol Complex Refractive Index and Size Distribution Variations on Extinction and Absorption Coefficients Part I: Tabulated Computational Results, ECOM Data Report ECOM-DR-77-5, Atmospheric Sciences Laboratory, US Army Electronics Command, White Sands Missile Range, NM. AD A049031

38. Patterson, E. M., D. A. Gillette, and G. W. Grams, 1976, "The Relation between Visibility and the Size-Number Distribution of Airborne Soil Particles," J. Appl. Meteorol., 15:470-478.
39. Bergstron, R. W., 1973, "Extinction and Absorption Coefficients of the Atmospheric Aerosol as a Function of Particle Size," Beit. Phys. Atmos., 46:223-234.
40. Waggoner, A. P., M. B. Baker, and R. J. Charlson, 1973, "Optical Absorption by Atmospheric Aerosols," Applied Optics, 12:896.
41. Jennings, S. G., 1977, "The Absorption Coefficient as a Function of Complex Index of Refraction and Particle Size Distribution," ECOM report in preparation.
42. Plass, G. N., 1966, "Mie Scattering and Absorption Cross Sections for Absorbing Particles," Applied Optics, 5:279-285.
43. Schleusener, S. A., J. D. Lindberg, K. O. White, R. L. Johnson, 1976, "Spectrophone Measurements of Infrared Laser Energy Absorption by Atmospheric Dust," Applied Optics, 15:2546-2550.
44. Gradshteyn, I. S., and I. M. Ryzhik, 1965, Table of Integrals, Series, and Products, Academic Press, New York and London, p. 307.

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